
The Chuacús Metamorphic Complex, central Guatemala: geochronological and geochemical constraints on its Paleozoic - Mesozoic evolution

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ABSTRACT

The Chuacús Metamorphic Complex is located in Central Guatemala, between the Polochic and Motagua fault zones. It is made up of complexly intercalated, mafic and felsic high-grade gneisses, amphibolites, pelitic and quartzofeldspathic metasediments and subordinate marbles. Mafic dikes and lenses metamorphosed to amphibolite and eclogite facies are tholeiitic and similar to mid-ocean ridge basalts. In contrast, metamorphosed intrusives (gabbro, diorite and granite) are calc-alkaline and have the geochemical signature of arc magmas. Laser Ablation Inductively Coupled Plasma Mass Spectrometry U-Pb zircon geochronology allows the recognition of three episodes of metamorphism. The first eclogite facies metamorphism (M1) is bracketed between Ordovician magmatism in the northern Chuacús Metamorphic Complex and the neighboring Rabinal granitic suite; the second corresponds to an Upper Triassic period of arc magmatism and migmatization (M2); the third high-grade metamorphic event (M3) occurred during the Late Cretaceous.

The tectonic evolution of the Chuacús Metamorphic Complex began during the Early Paleozoic as a basin in the Rheic Ocean that received detrital material from the Maya Block, Acatlán and southeastern México. The Chuacús Metamorphic Complex evolved to an active margin that subducted to HP conditions during the Mid-Late Paleozoic, and then was exhumed and involved in two tectonothermal events during the Upper Triassic and Late Cretaceous. The Chuacús Metamorphic Complex was accreted to the southern Maya Block during the Late Cretaceous, as a result of the convergent tectonics between the latter and either the Greater Antillean arc or the Chortís Block.

KEYWORDS | Chuacús Complex. Guatemala. U-Pb geochronology. Maya Block. Caribbean. Basement Tectonics.

INTRODUCTION

The tectonic limit between the North America and Caribbean plates is composed of a complex mosaic of crustal blocks, which are juxtaposed and limited by faults (e.g., Ortega-Gutiérrez et al., 2007; Rogers et al., 2007; Brueckner et al., 2009; Ratschbacher et al., 2009) (Fig. 1). Those crustal blocks range in age from the Proterozoic (e.g. Yoro area, central Honduras, Manton, 1996) to Early-Late Paleozoic and Mesozoic (Altos Cuchumatanes, Guatemala, Solari et al., 2009, 2010a; Maya Mountains, Belize, Martens et al., 2010; Rabinal suite, Guatemala, Ortega-Obregón et al., 2008; Chuacús Complex, Ortega-Gutiérrez et al., 2004; Ratschbacher et al., 2009; Las Ovejas Complex, Ratschbacher et al., 2009), and were finally amalgamated during the Late Mesozoic-Cenozoic (Harlow et al., 2004; Ortega-Gutiérrez et al., 2004; Martens et al., 2007; Brueckner et al., 2009; Ratschbacher et al., 2009). Recent works shed light on the tectonic evolution of the Motagua fault zone as the complex tectonic limit between the North America and Caribbean plates (Harlow et al., 2004; Brueckner et al., 2009; Pindell et al., 2006; Pindell and Kennan, 2009). The tectonic significance of the continental blocks involved in the suturing, however, is still a matter of debate, especially on the southernmost edge of the North America plate (i.e., the Maya Block). The classical view on the tectonic limit is represented by a suture along the Motagua fault, involving the Maya Block (North America plate) to the north, and the Chortís Block (Caribbean plate) to the south (e.g., Dengo, 1969). More recently, several works have added complexity to such a simple model, involving the collision of the Greater Antilles to the south of the Maya Block during the Late Cretaceous, prior to docking of the Chortís Block (e.g., Pindell and Barret, 1990; Harlow et al., 2004; Giunta et al., 2006; Brueckner et al., 2009; Pindell and Kennan, 2009; Ratschbacher et al., 2009). Further complications have arisen from the works of Ortega-Gutiérrez et al. (2007) and Ortega-Obregón et al. (2008), who suggested and argued for the existence of discrete crustal blocks or suspect terranes in the southern Maya Block, extending from central Guatemala across the Motagua suture zone to well inside the Chortís Block of southern Guatemala and Honduras (e.g. Rogers et al., 2007). One of the key localities to test the possible existence of those crustal entities, and one of the least studied, is the Sierra de Chuacús in central Guatemala. This paper presents new geochronological and geochemical data of crustal rocks in the Sierra de Chuacús in order to provide further constraints to its tectonic evolution and the ongoing discussions on its allochthonous vs. autochthonous nature and tectonic evolution.

PREVIOUS WORKS

The Sierra de Chuacús is located in central Guatemala, between the Polochic and Motagua faults (Fig.1 and 2). The

Chuacús Metamorphic Complex is made up of an intricate sequence of polydeformed, locally retrograded high-grade metamorphic rocks, both igneous and sedimentary in origin. Previous works report Silurian ages with Grenvillian protoliths in the northern sector of the Sierra de Chuacús (Gomberg et al., 1968). The presence of a Carboniferous high-grade metamorphic event affecting a Grenvillian protolith was identified by Ortega-Gutiérrez et al. (2004) in El Chol, in the middle of the Sierra de Chuacús. Martens et al. (2005, 2007) questioned the precision of that age and reported a Triassic age for similar rocks. A recent work by Ratschbacher et al. (2009) reported Proterozoic ages in the northern Sierra de Chuacús, with some discordant analyses yielding imprecise lower intercepts in the 638-477Ma range. Another sample they dated in the southern Sierra de Chuacús yielded Proterozoic and Devonian inheritance, as well as Triassic magmatic ages that postdate eclogite metamorphism and were later affected by high-temperature (HT) metamorphism during the Late Cretaceous. Petrological data suggest the existence of a HT eclogite facies event (~680-750°C, ~22-24kbar, Ortega-Gutiérrez et al., 2004; Martens et al., 2007; Ratschbacher et al., 2009), which cooled to about 580-600°C and 10-14kbar during decompression (Ortega-Gutiérrez et al., 2004; Ratschbacher et al., 2009). While there is still debate about the existence of one or two eclogite facies from metamorphic pulses affecting the Chuacús Metamorphic Complex; pre-Carboniferous according to Ortega-Gutiérrez et al. (2004), pre-Triassic according to Ratschbacher et al. (2009), or after the Triassic but before the Late Cretaceous according to Martens et al. (2007); these authors agree on Late Cretaceous cooling and exhumation (broad K-Ar, Ar and Rb-Sr age range of ~75 to ~66Ma, see also Ortega-Obregón et al., 2008).

Metamorphic and igneous rocks of comparable age and metamorphic grade are present elsewhere in southern México and central-north Guatemala. In the Acatlán Complex of southern México, basement of the Mixteco terrane, megacrystic granites have yielded ~480-440Ma magmatic ages, similar to the ages calculated for rift-related tholeiites in the same unit (Miller et al. 2007; Keppie et al. 2008a, b). They are intimately associated with high pressure metamorphism that occurred either during the Cambro-Ordovician (Vega-Granillo et al., 2007) or the Devonian-Mississippian and was associated with the closure of the Rheic Ocean (e.g. Nance et al., 2006, and references therein). In southeastern México, the Guichicovi complex is characterized by Grenvillian granulites (Weber and Köhler, 1999) intruded by Permian granites (La Mixtequita Massif, Torres et al., 1999). The Chiapas Massif crops out farther to the southeast, toward the Guatemala border. The massif is characterized by Ordovician S-type granites and amphibolites (Pompa-Mera et al., 2008), as well as widespread Late Permian granites that intrude medium to high-grade metasediments and orthogneisses, with Nd model

ages between 1.0 and 1.4Ga (Schaaf et al., 2002; Weber et al., 2006a, 2007, 2008). The metasediments of the Lower Santa Rosa Formation, which crop out in the Chiapas Massif, were recently dated by Weber et al. (2009), establishing the maximum depositional ages as defined by youngest zircons clustering around 315-325 and 331-341Ma. In Central Guatemala, just north of the Baja Verapaz Fault Zone, Ortega-Obregón et al. (2008) dated several peraluminous granites between 453 and 462Ma (lower intercepts of discordant data, U-Pb by Isotope-Dilution Thermal Ionization Mass Spectrometry (ID-TIMS)) and placed them at the very edge of what they considered to be the tectonic southern limit of the Maya Block. In the Maya Mountains of Belize, Martens et al. (2010) reported an Early Devonian Laser Ablation Inductively Coupled Plasma Mass Spectrometry (LA-ICPMS) age of 406Ma for rhyolites interbedded with conglomerates, which confirms the ~405-420Ma ages previously published by Steiner and Walker (1996).

GEOLOGY OF THE CHUACÚS METAMORPHIC COMPLEX

The Chuacús Metamorphic Complex is a polymetamorphic, originally high-grade and essentially metasedimentary formation, which bears textural and mineralogical evidence of at least one eclogite-facies event following a P-T decompressional path that could have departed near the field of ultrahigh pressure metamorphism (UHPM) and, in any case, representing continental subduction (e.g. Ortega-Gutiérrez et al., 2004; Martens, 2009). The Chuacús Metamorphic Complex is mainly composed of quartzofeldspathic, garnet and mica bearing gneisses; two micas, garnet, kyanite, rutile schists; minor marbles and calcsilicates and abundant orthogneisses, which range from mafic (garnet, hornblende, rutile and plagioclase) to granitic in composition. A large, Late Cretaceous shear zone, named Baja Verapaz Shear Zone by Ortega-Obregón et al. (2008) limits the Chuacús Metamorphic Complex in the north. The kinematics of the

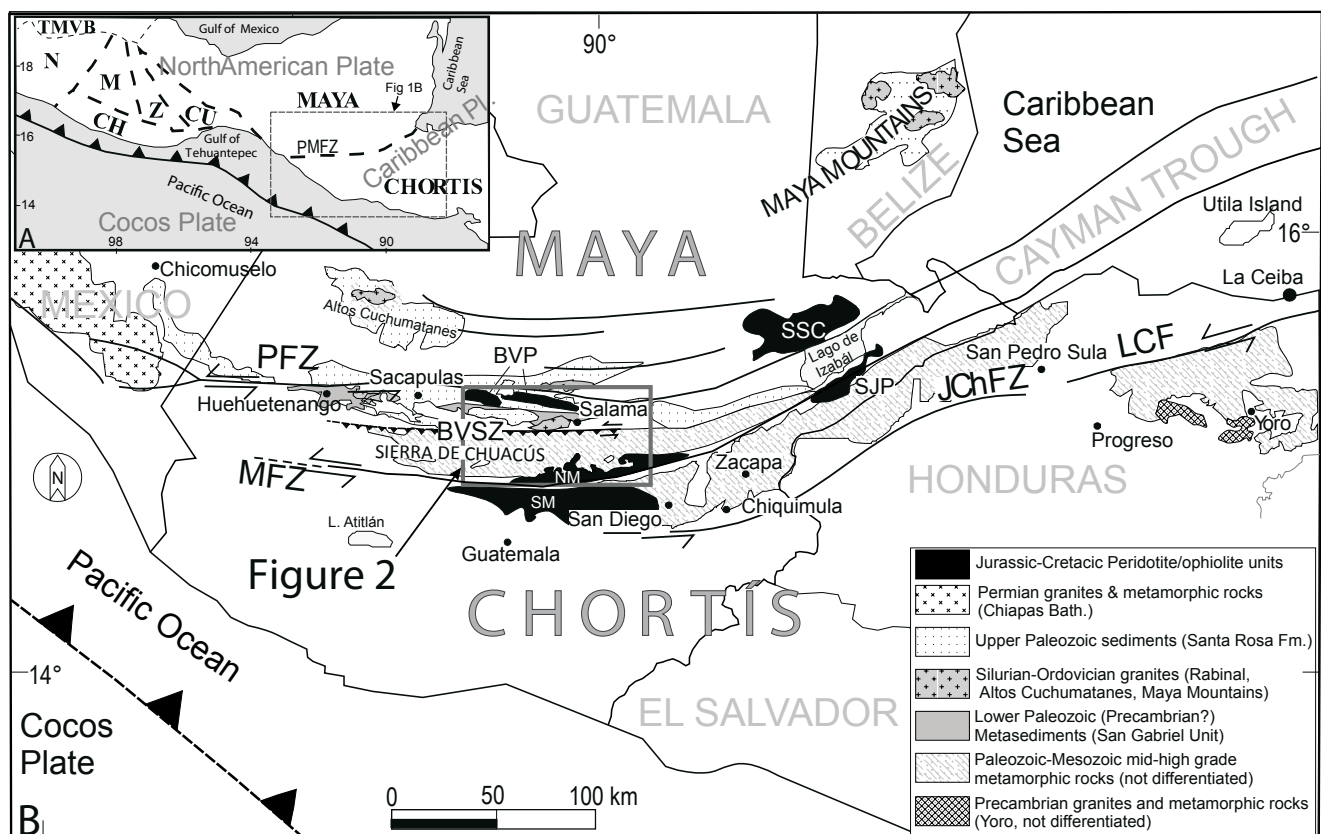


FIGURE 1 | A) Main tectonic subdivision of crustal blocks in southern Mexico and Central America. N: Nahuatl-Guerrero terrane; TMVB: Trans-Mexican Volcanic Belt; M: Mixteco terrane; CH: Chatino terrane; Z: Zapoteco terrane; CU: Cuicateco terrane; MAYA: Maya Block; CHORTIS: Chortís block. Nomenclature after Sedlock et al. (1993). PMFZ: main (simplified) trace of the Polochic-Motagua fault zone. B) Schematic geological map showing basement rocks of central Guatemala, southeasternmost Mexico, and northwestern Honduras. Geology is compiled from Kesler et al. (1970); Anderson et al. (1973); Steiner and Walker (1996); Weber et al. (2005, 2006); and unpublished data by the authors. PFZ: Polochic Fault Zone; MFZ: Motagua Fault Zone; BVFZ: Baja Verapaz Fault Zone; JChFZ: Jocotán-Chamelecón Fault Zone; LCF: La Ceiba Fault. Main ophiolitic bodies are shown in black: SSC: Sierra de Santa Cruz; BVP: Baja Verapaz; SJP: S. Juan de Paz; NM: north Motagua unit, ophiolite and mélange; SM: south Motagua unit, ophiolite and mélange. Modified from Ortega-Obregón et al. (2008).

Baja Verapaz Shear Zone indicates thrusting to the NNE with a minor left-lateral component. It mainly affected a sequence of low-grade pre-Ordovician slates and phyllites, informally named the San Gabriel sequence, as well as some igneous units belonging to the Ordovician Rabinal Granite Suite (Ortega-Obregón et al., 2008). At the southern edge of the Baja Verapaz Shear Zone the low metamorphic grade increases to amphibolite facies, where the Chuacús Metamorphic Complex actually begins. Chuacús Metamorphic Complex orthogneisses present a large age range and variable degrees of deformation. However, intrusive contacts are often visible in fresh outcrops. One of the Chuacús Metamorphic Complex exposures which deserves a more detailed description is the El Chol. At this location, banded mafic gneisses are intimately associated with metamorphosed granitic-trondhjemitic dikes. Local pods of granitic melt, slightly discordant to concordant with the main foliation, strongly suggest that the mafic bands are migmatitic restites, whereas the granitic veins are neosomatic leucosomes. It is important to underline that El Chol is the only locality in the Chuacús Metamorphic Complex where at least three deformation events are superposed and clearly distinguishable. Pegmatites are also widespread throughout the entire Chuacús Metamorphic

Complex, ranging from decimetric bands to more than several hundreds of meters wide. Eclogite is the most important lithology recognized in several outcrops. These high-pressure rocks are mainly present as deformed lenses in leucocratic and deformed gneisses, indicating a remnant of an important metamorphic event (cf. Ortega-Gutiérrez et al., 2004, for representative microprobe analyses).

The high-pressure metamorphism is widespread in the southern Sierra de Chuacús between El Chol and the northern Motagua allochthonous units, which constitute the tectonic southernmost limit of the Chuacús Metamorphic Complex. Eclogite-facies metamorphism is not recognized, however, in the western Chuacús (e.g. Kesler et al., 1970; Ratschbacher et al., 2009) or in the easternmost known outcrops, which are located in the southern Sierra de Las Minas (easternmost outcrops of Fig. 2).

MAIN PETROGRAPHIC FEATURES

The following petrological data summarize and complement previous work (Ortega-Gutiérrez et al., 2004) on some fundamental petrographic and textural

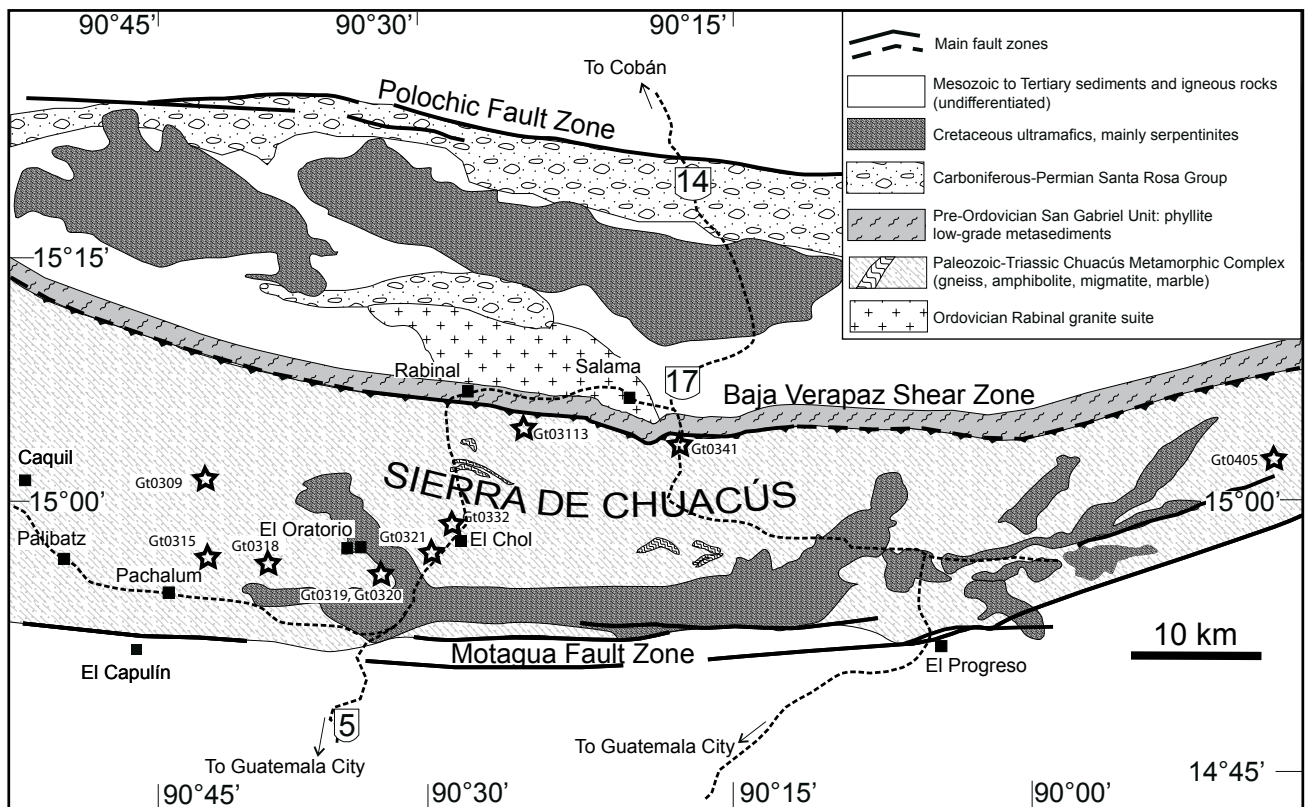


FIGURE 2 | Simplified geologic map of the Sierra de Chuacús area, Central Guatemala. Modified from Ortega-Gutiérrez et al. (2004). Sampled localities are indicated.

characteristics of the Chuacús Metamorphic Complex as studied at its type-area, the Sierra Chuacús between Granados and Salamá in central Guatemala. Based on detailed petrographic studies of representative samples, actual and relict evidence for high-pressure metamorphism may be found in most metapelitic, quartzofeldspathic, calcareous, and mafic lithologies of the Chuacús Complex in its type-area. The metapelites (see Table 1 for representative Wavelength Dispersive X-Ray Spectrometry (WDS) chemical analyses of the main aluminous phases) include the characteristic high-pressure assemblage kyanite - garnet - rutile - phengite - quartz, with minor phases in the matrix or as inclusions inside high-pyrope garnet ($\text{Alm}_{54}\text{Pyr}_{40.5}\text{Gr}_{4.2}\text{Spes}_{1.3}$) and kinked kyanite (Fig. 3A) such as quartz, phengite, ilmenite, titanomagnetite, zoisite, tourmaline, staurolite, phlogopite/biotite, chloritoid and monazite. Rutile inclusions in kyanite and garnet are particularly abundant and commonly show non-prismatic shapes (Fig. 3D). The high-pressure assemblages in the calcareous rocks include: a) high-Mg carbonate - rutile/titanite - quartz - white mica, with zircon, pyrrhotite and apatite as minor phases (marbles), and b) amphibole (pargasite/tremolite) - zoisite - quartz - clinopyroxene - white mica - clinocllore - serpentine - titanite (calcsilicates). Energy Dispersive X-Ray Spectrometry (EDS) analyses of high relief grains detected inclusions of corundum in some reddish pleochroic titanites, and lamellar rutile needles up to $32 \times 0.25 \mu\text{m}$ (Fig. 3B) that may be oriented in three different crystallographic planes in the cores of some zoisites. The composite association tremolite - zoisite - titanite/rutile (S1) overprinted by jadeite - lawsonite - albite - quartz - carbonate (S2) (see Fig. 3F) locally reflects two high-pressure events verified at contrasting higher and lower temperatures. Lawsonite and jadeite were not verified by chemical analyses, but both phases display their distinctive optical properties (Ortega-Gutiérrez et al., 2004).

Quartzo-feldspathic rocks are characterized by assemblages with quartz - garnet - albite - phengite - biotite - zoisite - rutile - subcalcic amphibole, with accessory potassium feldspar, carbonates and rare green clinopyroxene (probably omphacite) included in albite porphyroblasts, indicating relict high-pressure conditions. Common rounded zircons suggest an original sedimentary protolith. Mafic rocks contain hydrous and anhydrous eclogitic remnants with the typical assemblage omphacite - garnet - rutile - quartz accompanied by variable and commonly large amounts of amphibole (barroisite and taramite), phengite and zoisite/epidote. The eclogitic lithologies sometimes preserved radial cracking of garnet around quartz inclusions, as well as lamellar inclusions of rutile along 3-4 crystallographic directions and polycrystalline quartz - potassium feldspar inclusions.

TABLE 1 | Chuacús Metamorphic Complex metapelite Wavelength Dispersive X-Ray Spectrometry (WDS) mineral chemistry. Methodology described in Ortega-Gutiérrez et al., 2004

Oxide	St	St	St	St	Ky	Ky	Grt	Phe
SiO ₂	28.02	28.03	28.10	27.95	36.13	36.83	41.17	50.49
TiO ₂	0.46	0.43	0.42	0.37	0.00	0.00	0.00	0.12
Al ₂ O ₃	52.31	52.44	52.82	52.83	61.96	62.97	24.15	26.61
Cr ₂ O ₃	0.03	0.05	0.00	0.00	0.05	0.02	0.00	0.00
FeO	14.21	14.13	13.62	13.46	0.84	0.80	22.78	4.48
MnO	0.08	0.00	0.04	0.01	0.02	0.02	0.53	0.02
MgO	1.57	1.57	1.49	1.51	0.00	0.00	9.59	2.96
CaO	0.00	0.01	0.00	0.00	0.00	0.02	1.39	0.01
Na ₂ O	0.15	0.00	0.00	0.13	0.00	0.00	0.00	0.21
K ₂ O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	11.63
Total	96.83	96.66	96.49	96.26	99.00	100.66	99.61	96.53
Oxygens	46	46	46	46	5	5	12	11
Cations								
Si	7.88	7.88	7.89	7.86	0.99	0.92	3.08	3.39
Ti	0.10	0.09	0.09	0.08	-	-	-	0.01
Al	17.33	17.37	17.48	17.52	2.00	2.00	2.13	2.10
Cr	0.02	0.03	-	-	0.01	-	-	-
Fe	3.34	3.32	3.20	3.17	0.02	0.02	1.43	0.25
Mn	0.02	-	0.01	0.00	-	-	0.03	-
Mg	0.66	0.66	0.62	0.63	-	-	1.07	0.30
Ca	-	0.00	-	-	-	0.00	0.11	-
Na	0.08	-	-	-	-	-	-	0.03
K	-	-	-	-	-	-	-	0.99
Sum	29.42	29.35	29.29	29.26	3.02	2.94	7.85	7.07

St: staurolite; Ky: kyanite; Grt: garnet; Phe: phengite

The polymetamorphic evolution of the Chuacús Metamorphic Complex is evident in the wide variety of reaction (prograde and retrograde) textures that characterize most rocks, particularly the eclogitic units. Prograde coronas include garnet on omphacite (Fig. 3E), or on biotite, plagioclase, chlorite and apatite, as well as occasional rims of clinopyroxene on garnet followed by amphibole - plagioclase symplectites (Fig. 3C) that suggest alternating metamorphic episodes of hydration and dehydration. Retrograde or decompression coronitic textures are characterized by symplectites of amphibole-plagioclase and clinopyroxene - albite on omphacite, subcalcic amphibole and garnet, epidote on garnet, apatite on allanite, biotite on phengite, albite on white mica and garnet, titanite on ilmenite and rutile, epidote on zoisite, green on brown biotite, and chlorite on garnet.

ANALYTICAL METHODS

Geochemistry

Major and trace element analyses were performed at Actlabs (Ancaster, Canada). Major elements were determined by X-ray fluorescence, whereas trace elements were measured by ICP-MS employing lithium metaborate/tetraborate fusions. Sm-Nd isotopic ratios were measured by Thermal Ionization Mass Spectrometry (TIMS) at the Laboratorio Universitario de Geoquímica Isotópica of the Universidad Nacional Autónoma de México (UNAM), using a Finnigan MAT 262 system equipped with eight

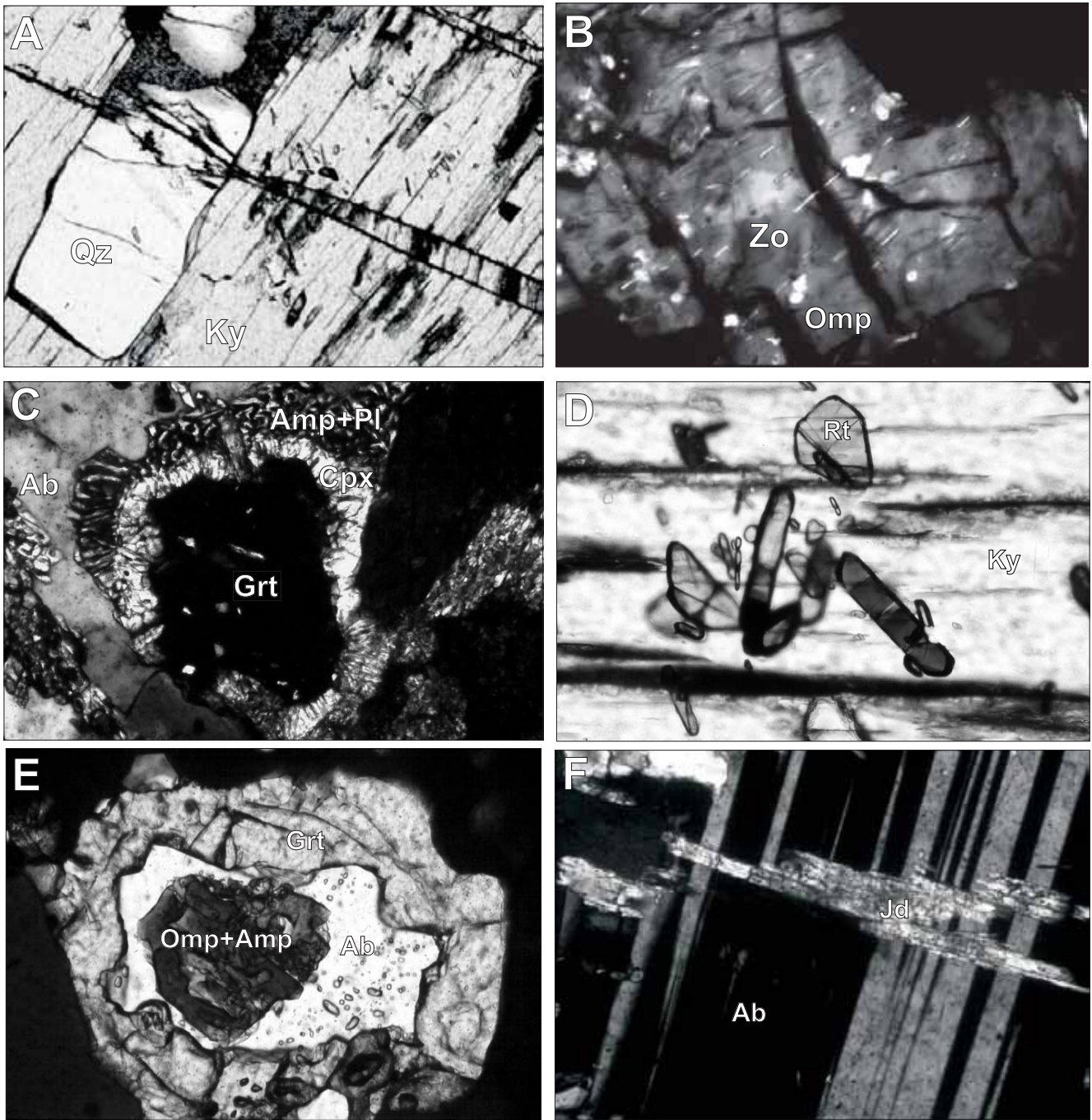


FIGURE 3 | Some representative photomicrographs of petrographic features described in the text. **A)** Kink bands nucleated on an elongate quartz inclusion in kyanite. Plane polarized light, field of view is 560 μ m. **B)** Lamellar inclusions of rutile (bright needles) in zoisite in a high-pressure calc-silicate. Crossed polarized light, field of view is 220 μ m. **C)** Multiple coronas formed by retrogression of mafic eclogites; garnet appears rimmed by clinopyroxene, then by a symplectite of amphibole-plagioclase, and finally by an albite-rich matrix. Crossed polarized light, field of view is 2,240 μ m. **D)** Rutile inclusions in kyanite, many crystals show a peculiar “butterfly” shape and may be reddened by exsolution of hematite. Crossed polarized light, field of view is 220 μ m. **E)** Composite coronitic texture formed by a central zone of clinopyroxene-amphibole symplectite, rimmed by a wide band of albite followed by a band of garnet with rutile/titanite inclusions; this garnet forms a large inclusion in subcalcic amphibole. Plane polarized light, field of view is 2,240 μ m. **F)** Euhedral jadeite (optically determined by extinction angle, birefringence and relief) grown across an albite porphyroblast. Crossed polarized light, field of view is 220 μ m.

Faraday cups. Sample preparation and measurement procedures for isotopic analyses have been described by Schaaf et al. (2005). $^{143}\text{Nd}/^{144}\text{Nd}$ ratios were normalized to $^{146}\text{Nd}/^{144}\text{Nd}=0.72190$ and corrected to a La Jolla standard value of $^{143}\text{Nd}/^{144}\text{Nd}=0.511860$. The long term reproducibility of La Jolla standard at Laboratorio Universitario de Geoquímica Isotópica ($^{143}\text{Nd}/^{144}\text{Nd}$) is 0.511875 ± 22 (2σ , $n=152$). Geochemical results are reported in Table 2.

Geochronology

The samples selected for geochronology, (3 to 5kg) were crushed and the minerals separated following common techniques such as jaw crushing, milling, Wifley separation, Frantz magnetic separation, and heavy liquids (see also Solari et al., 2007). The zircon concentrates were carefully observed under a binocular microscope. Between 500 and 150 crystals were carefully selected from each sample, in order to obtain all the variations in morphology, shape and color. They were mounted in epoxy resin, ground to expose roughly half of the crystal, and then imaged by cathodoluminescence using an ELM 3R luminescope (Marshall, 1988).

LA-ICPMS U-Pb analyses were performed at Laboratorio de Estudios Isotópicos (LEI), Centro de Geociencias, UNAM, using a Resonetics "Resolution M50" 193nm excimer laser workstation coupled with Thermo Xii Series quadrupole mass spectrometry. Details of the analytical setup employed during this study are described in Solari et al. (2010b). 25s of gas background acquisition was followed by 30s of ablation, carried out in He atmosphere, employing 160mj laser energy, corresponding to an on-target fluence of $8\text{J}/\text{cm}^2$, $34\mu\text{m}$ diameter spot and a repetition rate of 5Hz, generating a drill rate of $\sim 0.7\mu\text{m}/\text{s}$. The background average was subtracted from net intensities measured for each isotope. Acquisition involved an alternance of 2 analyses of Plešovice reference zircon ($\sim 337\text{Ma}$, Sláma et al., 2008), NIST 610 standard glasses and 5 unknown zircons, using standard-unknown bracketing method (e.g. Solari et al., 2010b) to allow down-hole fractionation corrections to be performed using an in-house developed software (Tanner and Solari, 2009). NIST 610 glass standard is used to recalculate the concentrations of interest, for instance U and Th, by normalizing such elements with ^{29}Si . Other isotopes are observed during analysis; such as P, Ti, REE; to monitor the presence of inclusions (and in due case discard it) and to produce zircon REE patterns, which can be helpful to interpret the calculated ages. REE zircon patterns were normalized to chondrite values of McDonough and Sun (1995). Recalculated values are available in Table I (Electronic Appendix in www.geologica-acta.com). An estimation of the zircon crystallization temperature is calculated by the Ti content, according to the Ti-in thermometer of Watson et

al. (2006). Precision on measured $^{207}\text{Pb}/^{206}\text{Pb}$, $^{206}\text{Pb}/^{238}\text{U}$ and $^{208}\text{Pb}/^{232}\text{Th}$ ratios typically was $\sim 0.7\%$ 1σ relative standard deviation. Replicate analyses of the Plešovice zircon indicate an external reproducibility of 0.8%, 0.7% and 1.6% on the measured $^{207}\text{Pb}/^{206}\text{Pb}$, $^{206}\text{Pb}/^{238}\text{U}$ and $^{208}\text{Pb}/^{232}\text{Th}$ ratios, respectively. These errors are quadratically included in the quoted uncertainties for individual analyses of the analyzed zircons. ^{204}Pb is not monitored during these analyses, because its signal is swamped by the ^{204}Hg contained in the carrier gases. Common Pb correction was thus performed employing the algebraic method of Andersen (2002). A filter was then applied to ensure the quality of selected analyses, consisting in evaluation of the concordance. For grains with ages of $<1000\text{Ma}$, the analysis was considered concordant if the $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{235}\text{U}$ ages differed by less than 10%. For the grains with ages $>1000\text{Ma}$, the same test was carried out considering $^{206}\text{Pb}/^{238}\text{U}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ ages. The concordia, probability density distribution and histogram plots, as well as age-error calculations are performed using Isoplot v. 3.70 (Ludwig, 2008). The Tuff-Zirc algorithm (Ludwig and Mundil, 2002) contained in the same software was used to calculate the mean $^{206}\text{Pb}/^{238}\text{U}$ ages and their errors, as well as to filter for outliers, which are preferred for grains younger than 1000Ma. Details of analyzed zircons, for each samples, are reported in Table II of the Electronic Appendix (www.geologica-acta.com).

RESULTS

Geochemistry

Most of the samples studied here underwent high-grade metamorphism, and therefore geochemical classifications and interpretations must be taken with caution. Nevertheless, comparisons of relatively immobile element concentrations and ratios of the studied rocks with those found in unaltered volcanic rocks can provide insights into the sources and processes that lead to their formation.

In terms of geochemistry and petrography, rocks from the Chuacús Metamorphic Complex can be divided in three different groups: tholeiitic metabasalts (amphibolites), calc-alkaline orthogneisses and pelitic gneisses. Migmatitic gneisses are also present, but they will be described separately for simplicity.

Amphibolites are found as metamorphosed lenses, or deformed dikes in orthogneisses and migmatites (e.g. samples Gt0322, Gt0327, Gt0329, Gt0334 in Table 2). They are fairly homogeneous in terms of their major elements, and show enrichment in Fe_2O_3 and TiO_2 when compared to the rest of the sequences (AFM diagram of Fig. 4A). Even though the MgO contents are not what would be expected

TABLE 2 | Representative analyses of serpentine from antigorite serpentinites (normalized to 10 and 8 OH)

SAMPLE	Chuacús Metamorphic Complex																
	G10303 Arc orthogneiss	G10309 Arc orthogneiss	G10314 Metasediment	G10320 Arc orthogneiss	G10321 Arc orthogneiss	G10322 MORB	G10327 MORB	G10329 MORB	G10333 Migmatite Leucosome	G10334 Migmatite Melanosome	G10341 Rabinal	GT0340 Rabinal	G103113 Rabinal	EC01 MORB	G10360 Metasediment	G10394 MORB	G103117 MORB
SiO ₂	50.72	74.81	71.91	46.78	69.43	48.15	47.65	49.60	73.02	46.55	64.35	73.45	75.84	48.44	52.25	48.13	47.56
Al ₂ O ₃	16.05	13.45	13.81	12.04	14.90	13.78	13.76	13.19	15.34	15.47	15.45	13.05	13.17	13.55	22.65	13.91	14.29
Fe ₂ O ₃	9.84	2.46	2.88	9.47	0.60	12.82	0.354	10.89	1.13	13.61	4.05	1.87	1.00	13.24	9.30	14.12	12.68
MnO	0.159	0.019	0.023	0.175	0.060	0.284	0.354	0.237	0.012	0.199	0.067	0.062	0.005	0.263	0.122	0.178	0.129
MgO	6.71	0.29	0.69	4.54	14.10	6.76	6.84	6.82	0.32	6.33	1.27	0.34	0.03	7.03	2.32	7.09	8.15
CaO	5.80	0.76	1.74	7.60	1.88	9.61	9.23	10.09	2.08	9.19	3.86	0.83	0.25	10.35	0.34	10.49	10.90
Na ₂ O	5.34	5.12	5.00	2.86	4.67	4.05	3.69	3.26	5.38	3.76	5.20	3.45	3.88	3.29	0.77	2.57	2.50
K ₂ O	1.42	2.19	1.98	0.62	3.66	1.14	0.66	2.80	1.50	1.24	2.04	4.97	5.51	0.15	5.48	0.29	0.23
TiO ₂	1.828	0.280	0.582	1.231	0.463	2.300	2.594	1.560	0.132	2.507	0.723	0.271	0.026	1.831	1.022	1.470	1.322
P ₂ O ₅	0.63	0.08	0.19	0.43	0.25	0.32	0.30	0.14	0.04	0.32	0.26	0.07	0.02	0.17	0.18	0.10	0.11
LOI	1.86	0.80	1.51	1.42	1.01	0.82	0.69	1.71	0.82	0.62	2.74	0.58	0.35	0.65	4.36	1.27	2.23
TOTAL	100.36	100.26	100.32	100.23	99.87	99.97	100.31	100.30	99.78	100.00	100.01	98.95	100.08	99.07	98.79	99.60	100.11
Sc	20	2	9	40	5	41	42	36	2	25	11	3	1	47	26	42	42
Be	2	2	1	2	2	2	2	2	1	1	1	2	2	1	5	BDL	BDL
V	134	13	27	172	28	381	370	270	13	245	65	18	5	423	104	337	325
Cr	111	BDL	BDL	1.020	BDL	213	122	189	BDL	67	BDL	BDL	50	209	79	173	251
Co	33	3	6	54	4	40	39	38	2	46	9	2	BDL	44	20	47	48
Ni	91	BDL	BDL	326	BDL	85	100	124	BDL	112	BDL	BDL	BDL	103	42	102	113
Cu	107	BDL	BDL	117	BDL	70	85	64	10	30	53	BDL	BDL	38	BDL	73	183
Zn	107	BDL	BDL	97	42	174	559	134	BDL	133	43	BDL	BDL	95	173	118	118
Ga	18	21	18	13	16	19	17	17	16	19	16	16	19	19	33	18	18
Ge	1.8	1.0	0.7	1.3	1.2	1.4	1.7	1.2	0.5	0.9	1.0	1.5	1.0	1.5	2.6	2.1	2.2
Rb	24	103	39	3	66	24	4	61	22	15	66	176	123	2	212	2	5
Sr	628	98	243	214	224	76	112	65	774	139	641	124	59	105	128	360	225
Y	36.3	17.1	31.4	28.7	25.2	49.6	59.9	40.6	1.1	23.5	15.5	27.9	11.7	50.9	76.5	22.2	22.2
Zr	237	210	436	126	230	156	160	83	4.6	158	321	180	85	141	157	80	75
Nb	8.6	6.3	13.7	13.6	12.0	10.8	8.2	8.8	4.6	6.4	14.5	28.1	6.4	3.2	26.9	7.3	5.6
Cs	0.3	0.2	0.2	-0.1	0.5	0.2	-0.1	0.6	0.2	0.2	0.7	1.2	BDL	BDL	2.4	BDL	BDL
Ba	548	379	745	29	750	189	25	686	1.050	270	623	600	457	57	795	10	38
La	34	40.3	95.2	14.5	44.6	6.51	7.15	2.97	4.63	18.0	119	48.2	8.69	8.71	125	5.90	5.64
Ce	64.3	65.3	175	30.4	77.9	17.7	20.3	8.60	9.07	36.6	198	86.3	14.8	17.6	234	15.1	14.2
Pr	8.51	7.40	21.2	3.93	8.69	2.96	3.47	1.47	1.10	4.71	21.6	8.85	2.15	3.38	29.0	2.27	2.12
Nd	35.1	26.6	77.9	16.3	30.4	15.2	18.9	7.88	4.17	19.9	72.2	29.4	8.18	18.4	108	11.4	10.7
Sm	8.02	5.59	15.0	3.91	5.45	5.34	6.75	3.10	0.71	4.81	10.6	5.16	1.70	5.62	20.6	3.67	3.38
Eu	2.60	1.40	3.39	1.75	1.35	1.92	2.39	1.03	0.239	1.50	2.11	0.819	0.277	1.83	3.34	1.36	1.28
Gd	7.38	4.74	11.2	5.25	3.78	7.12	8.80	4.48	0.43	4.87	7.09	3.79	1.61	7.43	17.2	4.32	3.99
Tb	1.09	0.68	1.44	0.81	0.62	1.25	1.58	0.90	0.04	0.78	0.74	0.76	0.29	1.40	2.77	0.82	0.78
Dy	6.48	3.85	7.57	4.88	3.74	8.53	10.8	6.52	0.21	5.22	3.60	4.56	1.72	8.99	14.9	4.68	4.34
Ho	1.33	0.73	1.33	1.03	0.87	1.79	2.24	1.46	0.04	1.03	0.57	0.92	0.37	1.93	2.90	0.92	0.87
Er	3.52	1.81	3.11	2.80	2.51	5.20	6.43	4.26	0.10	2.92	1.37	2.94	1.22	5.94	9.07	2.76	2.58
Tm	0.532	0.245	0.404	0.428	0.414	0.847	1.02	0.708	0.014	0.448	0.159	0.549	0.196	0.884	1.35	0.368	0.372
Yb	3.46	1.48	2.32	2.72	2.72	5.28	6.38	4.38	0.11	2.67	1.10	3.40	1.34	5.53	8.02	2.38	2.27
Lu	0.513	0.210	0.310	0.288	0.432	0.785	0.951	0.676	0.022	0.389	0.173	0.533	0.223	0.811	1.21	0.358	0.329
Hf	5.5	6.6	11.8	2.8	5.9	4.6	5.0	2.8	2.5	4.3	8.3	5.5	3.4	4.1	4.6	2.4	2.5
Ta	0.56	0.10	0.48	0.48	0.96	0.26	0.25	0.28	BDL	0.90	0.40	2.99	0.63	0.12	1.25	0.25	0.21
W	0.8	BDL	BDL	BDL	BDL	0.6	0.6	BDL	BDL	0.9	BDL	0.6	0.8	BDL	BDL	BDL	BDL
Tl	0.14	0.47	0.18	0.21	0.29	0.13	0.08	0.26	0.09	0.06	0.28	0.87	1.04	0.07	2.12	BDL	0.10
Pb	9	BDL	BDL	BDL	8	BDL	BDL	BDL	10	BDL	12	13	7	BDL	8	BDL	BDL
Bi	BDL	BDL	BDL	BDL	BDL	BDL	BDL	BDL	-0.1	BDL	BDL	BDL	3.0	0.5	BDL	BDL	BDL
Th	3.08	2.81	2.56	3.05	6.15	0.38	0.39	0.20	0.34	1.08	23.1	20.8	4.91	0.26	29.0	0.43	0.42
U	0.73	1.35	0.67	0.69	1.35	0.31	0.44	0.21	0.15	0.25	2.56	4.68	2.64	0.10	3.52	0.12	0.12
¹⁴⁷ Sm/ ¹⁴⁴ Nd	0.125	0.115	0.103	0.132	0.107	0.196	0.202	0.221	0.092	0.139	0.062	0.104	-	-	-	-	-
¹⁴³ Nd/ ¹⁴⁴ Nd	0.512519 (19)	0.512033 (19)	0.511985 (18)	0.512708 (20)	0.512406 (20)	0.513093 (16)	0.513055 (19)	0.512779 (19)	0.512206 (18)	0.512626 (19)	0.512166 (12)	0.512245 (18)	-	-	-	-	-
eNd	-2.32	-11.80	-12.74	1.37	-4.53	8.88	8.13	2.75	-8.43	-0.23	-9.21	-7.67	-	-	-	-	-

Major elements expressed as wt%. Trace elements expressed as ppm

-: not determined

BDL: below than detection limits

*: last two digits in parentheses in the ¹⁴³Nd/¹⁴⁴Nd ratios refer to the 2 sigma absolute error

for pristine mantle-derived melts (Langmuir and Hanson, 1980), they have compositions that are typical of slightly differentiated tholeiitic basalts. Despite major element homogeneity, the REE patterns of these rocks are variable. Three samples have the typically depleted LREE and flat HREE pattern of normal mid-ocean ridge basalts

(N-MORB) and display small negative Eu anomalies, whereas two samples have relatively flat LREE and depleted HREE without Eu anomalies (Fig. 5A). Extended incompatible trace element patterns are also similar to MORB, although fluid mobile elements (Cs, Rb, Ba, U) have been likely perturbed by weathering since they resemble the enrichments that are often observed in altered-oceanic crust (Staudigel et al., 1996; Fig 5A). In

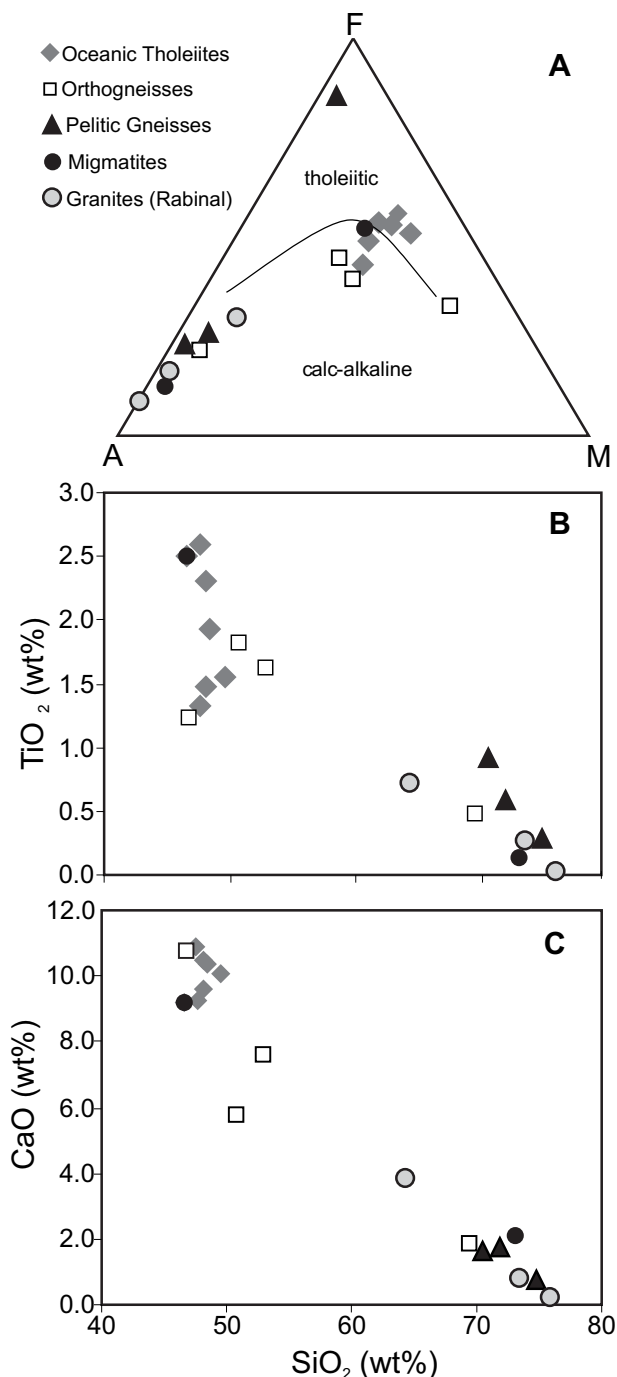


FIGURE 4 | Major element compositions of rock suites from Chuacús Complex. A) AFM diagram after Irvine and Baragar (1971). B) TiO₂ vs SiO₂. C) CaO vs SiO₂.

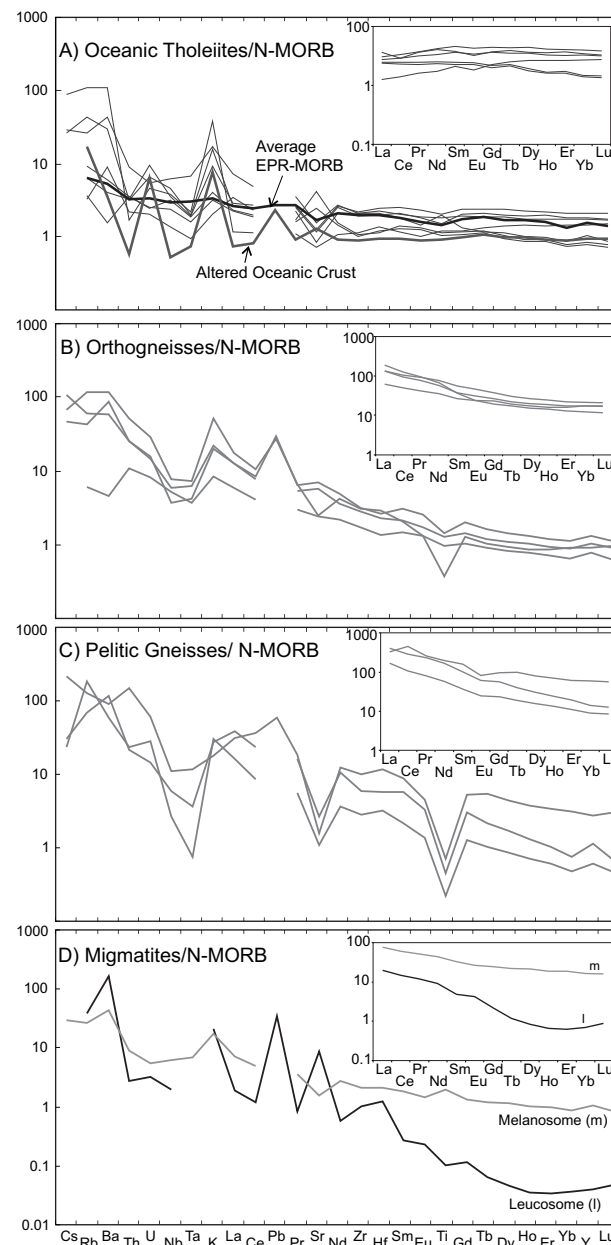


FIGURE 5 | Trace element geochemical composition of rock suites from Chuacús Complex, normalized to the N-MORB values of Sun and McDonough (1989). REE elements normalized to the chondrite values of McDonough and Sun (1995). Also shown for comparison is the composition of an average EPR-MORB (Lehnert et al., 2000), and of a typically altered oceanic crust (Staudigel et al., 1996).

fluid immobile elemental ratios (Th/Nd, La/Nd, Sr/Nd, Nd/Yb), these rocks display end-member compositions with values that overlap a typical Pacific MORB. Nonetheless, Nb/Ta ratios (up to 40) are much higher than in MORB, probably suggesting that an element like Ta did not remain completely immobile during high-grade metamorphism.

Orthogneisses (e.g., samples Gt0303, Gt0319, Gt0320, Gt0321 in Table 2) have variable compositions that range from gabbroic to granitic and display a systematic decrease in MgO, F₂O₃, and TiO₂ with increasing SiO₂, Na₂O and K₂O, characteristics that are typical of calc-alkaline magmas worldwide (Fig. 4A). Chondrite-normalized REE patterns show enrichment in LREE and slightly depleted HREE that are in contrast to what is observed in the tholeiitic amphibolites (Fig. 5B). A small negative Eu anomaly is only observed in a relatively evolved granite. Extended N-MORB normalized trace element patterns also display features that are typical of

arc magmas, such as Large Ion Lithophile Elements (LILE) enrichment relative to the High Field Strength Elements (HFSE). It is possible that fluid mobile elements like Cs, Rb, and Ba were perturbed by weathering and/or metamorphism; but high LILE/High Field Strength Elements ratios are also coupled with relative enrichment in Th (i.e. high Th/Nb ratios), which are considered as fluid immobile. This suggests that the “arc-like” geochemical signatures of these magmas were likely acquired by crustal inputs, either in the form of subducted sediments or by crustal contamination during ascent. Yet the Th/Nb ratios of the least evolved calc-alkaline orthogneisses are higher than those of similarly evolved tholeiitic metabasalts, strongly suggesting a primary origin in a magmatic arc environment.

Pelitic gneisses are silica and potassium-rich and have overall chemical characteristics that are typical of hemipelagic sedimentation (e.g. samples Gt0309, Gt0314, Gt0360 in Table 2). They are LREE enriched, with either flat or depleted HREE, and usually display negative Eu anomalies. Incompatible trace element patterns also display mature upper crustal features, with highly enriched LILE/High Field Strength Elements ratios that are coupled with negative Sr and TiO₂ anomalies (Fig. 5C). Interestingly, Nb/Ta ratios are also unusually high for upper crustal lithologies (up to ~60), probably indicating a preferential mobilization of Ta during metamorphism.

Distinct magmatic phases from El Chol migmatites have different compositions. A sample taken from the melanosome (Gt0334, Table 2) has the chemical character of an enriched oceanic basalt, with relatively high Nb-Ta contents and fractionated REE patterns. In contrast, the leucosome (Gt0333) is silica-rich and trondhjemitic (K₂O/Na₂O=0.3). The incompatible trace element pattern of the leucosome displays strong positive anomalies in LILE, K₂O, Pb and Sr that are not accompanied by equivalent enrichments in elements like Th and U. Y and the HREE are strongly depleted when compared to the LREE, but the trondhjemitic leucosome also displays a distinctive concave upward pattern in which elements like Ho and Er have the lowest relative concentrations (Fig. 5D). Also noteworthy is that Zr and Hf are enriched when compared to their nearest neighbors, and form a distinctive positive anomaly. These chemical characteristics are indicative of melting under garnet amphibolite facies conditions (Foley et al., 2002).

Nd isotopic compositions were recalculated to 225Ma in accordance to the most important magmatic and recrystallization event recorded by U-Th/Pb geochronology (see below). All amphibolites have positive εNd (up to ~10) values that are close to those of the calculated depleted upper mantle for that age. In contrast, orthogneisses are less radiogenic (εNd=3.23 to -1.95), with isotopic values that correlate negatively

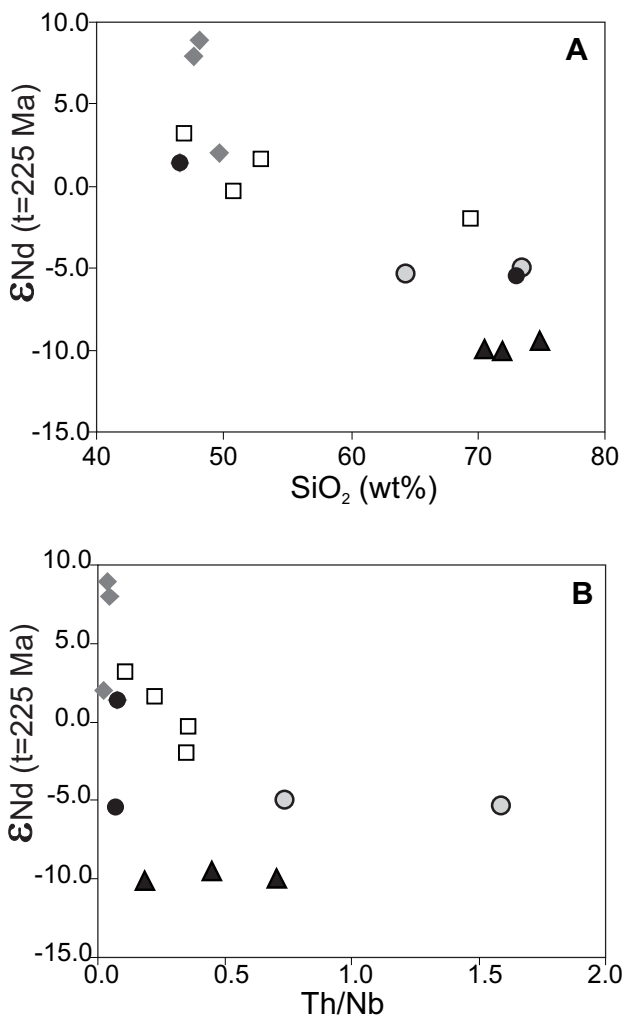


FIGURE 6 | Initial Nd isotopic compositions calculated to 225Ma for the Chuacús Metamorphic Complex correlate inversely with A) SiO₂ contents and B) Th/Nb ratios.

with SiO_2 and Th/Nb ratios (Fig. 6A, B). As it is commonly observed in ancient upper crustal lithologies, pelitic gneisses have very negative ϵNd ($\epsilon\text{Nd}=-10$), with values that are even less radiogenic than the Ordovician Rabinal Granite. Interestingly, the melanosome at El Chol has a positive ϵNd whereas the trondhjemitic leucosome displays a very negative value, indicating that both components are genetically unrelated and that additional crustal lithologies underwent partial melting.

Geochronology

Gt 0315

Sample Gt0315 is a pegmatite which intrudes a gabbroic boudin neck in the locality of Quebrada Honda. Zircons

separated from this pegmatite are colorless to pale pink, quite large (up to $280\mu\text{m}$ in length) and range in shape from large prismatic to stubby. Under cathodoluminescence they are mostly unzoned and characterized by low luminescence, although some show a faint sector-zoned to patchy-zoned cores, sometimes surrounded by moderately luminescent, faintly oscillatory-zoned rims (examples in Fig. 7). A total of 30 analyses were performed on 30 zircons belonging to this sample. The cores have an average age of $982\pm 12\text{Ma}$ (Fig. 8A), as well as high Th/U ratios and typical REE anomalies indicative of igneous origin (e.g., Rubatto, 2002 and Fig. 9A). The overgrowths, as well as those zircons which show a uniform cathodoluminescence, have very low Th/U ratios of less than 0.01, and show LREE depletion compared to HREE (Fig. 9B). The mean $^{206}\text{Pb}/^{238}\text{U}$ age of $74\pm 3/-1\text{Ma}$ obtained on such overgrowths



FIGURE 7 | Cathodoluminescence image of some representative zircons analyzed in this work. The circles represent the $34\mu\text{m}$ LA-ICPMS analyzed spot.

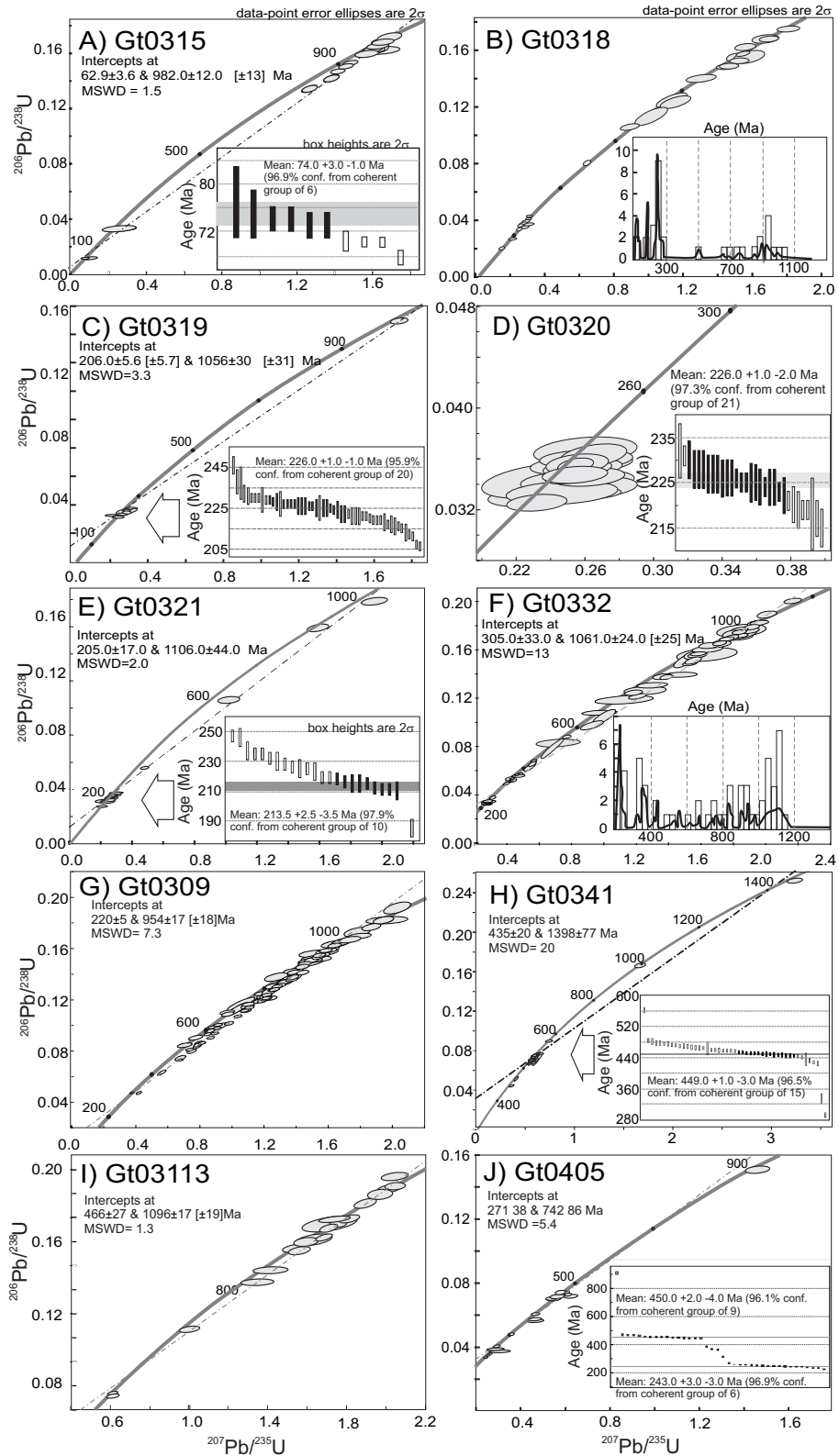


FIGURE 8 | U-Pb concordia diagrams of the samples dated by LA-ICPMS. Concordia, ages and errors are calculated using Isoplot 3.7 (Ludwig, 2008). The mean $^{206}\text{Pb}/^{238}\text{U}$ ages in the insets, 2-sigma errors and outlier filtering are calculated with the *TuffZirc* algorithm of Ludwig and Mundil (2002). For those samples containing abundant detrital zircons, i.e. A, C, G, and I, the inset diagrams are combined histograms and probability density plots. Preferred ages are common-Pb corrected using the algebraic method of Andersen (2002). $^{206}\text{Pb}/^{238}\text{U}$ ages are used if $<1000\text{Ma}$, whereas $^{207}\text{Pb}/^{235}\text{U}$ ages are used if $>1000\text{Ma}$. See text for further explanations.

is interpreted as indicating the recrystallization under high-grade metamorphic conditions (concordia of Fig. 8A).

Gt 0318

This sample is a high-grade metasedimentary rock cropping out in Los Altos, northeast of Pachalum (Fig. 2). It is made up of garnet, relic omphacite, white mica, hornblende and interstitial plagioclase. Rutile is present as an accessory mineral, often surrounded by titanite rims. We performed 45 U-Pb analyses on 45 zircons separated from this sample, 9 of which were discarded due to their large discordance. Zircons are in general elongated to ovoidal in shape, with polished terminations typical of sedimentary transport. Under cathodoluminescence they show complex zoning, with presence of inherited cores surrounded by variably luminescent overgrowths (examples in Fig. 7). The dated zircons yield a main cluster at 245Ma followed by other minor ones at 182, 895, 930, and concordant zircons at 125, 505, and several others between 606 and 1007Ma (concordia of Fig. 8B). Although the main provenance is thus Triassic, the few Jurassic to Cretaceous zircons indicate a young episode of sedimentation in the southern Sierra de Chuacús, postdated by high-grade metamorphism.

Gt 0319

This is a foliated gabbro cropping out in Los Limones locality, principally composed of clinopyroxene, garnet, plagioclase, minor biotite, titanite and opaque minerals. We dated 60 crystals by LA-ICPMS, 6 of which were discarded because their discordance was more than 10%. Under cathodoluminescence, the imaged zircons show a predominance of oscillatory (i.e., igneous) growth, with only minor bright overgrowths developed on some of the crystals (Fig. 7). Both oscillatory-zoned cores, as well as overgrowths, show an almost indistinguishable age pattern with a mean $^{206}\text{Pb}/^{238}\text{U}$ age of $226\pm 1\text{Ma}$. This age is interpreted as indicating the crystallization of the gabbroic protolith (concordia of Fig. 8C).

Gt 0320

Gt 0320 is an intermediate igneous rock composed of plagioclase, clinopyroxene, hornblende, opaque minerals, biotite, titanite, cropping out nearby sample Gt 0319. Cathodoluminescence shows the ubiquitous presence of oscillatory zoning, with eventual thin, high-luminescent rims developed around cores (Fig. 7). We dated 41 crystals belonging to this sample. Five analyses were omitted because of having >10% discordance. There is no inheritance in any of the dated crystals and therefore the mean $^{206}\text{Pb}/^{238}\text{U}$ age of $226\pm 1/-2\text{Ma}$ is interpreted as an indication the crystallization of this sample (Fig. 8D).

Gt 0321

This is a granodioritic orthogneiss, intensely foliated, cropping out in the Barranca Agua Caliente. Leucocratic bands, such as the one dated, contain mafic pods up to 30cm in size, which were recognized as retrogressed eclogites (cf. Ortega-Gutiérrez et al., 2004). Zircons separated from this sample are up to $350\mu\text{m}$ in size and pale yellow to dark red in color. Under cathodoluminescence, they show a frequent presence of oscillatory-zoned cores, surrounded by variably developed zones of oscillatory overgrowths. Of the 50 analyses which were performed on 50 zircons belonging to this sample, 17 were discarded because they were highly (17-42%) discordant, probably as a consequence of partial Pb loss. The other analyses yield a poorly defined Grenvillian upper intercept ($1106\pm 44\text{Ma}$) and a cluster of concordant to slightly discordant analyses corresponding to the lower intercept (Fig. 8E). Such analyses, showing an igneous Th/U ratio well above 0.05, yield a $^{206}\text{Pb}/^{238}\text{U}$ weighed mean age of $213.5\pm 2.5/-3.5\text{Ma}$, which is regarded as the best approximation of the crystallization age of this granodiorite.

Gt 0332

This sample belongs to a leucosome migmatitic orthogneiss of El Chol (Fig. 2). The poly-deformed leucocratic bands are interlayered, and some still preserve relict eclogitic assemblages. Previous U-Pb TIMS dating of zircons belonging to this sample yielded discordant results, with a lower intercept of $302\pm 52\text{Ma}$ and an upper intercept of $1024\pm 78\text{Ma}$ (Ortega-Gutiérrez et al., 2004). The lower intercept was interpreted as the age of migmatization.

The zircons selected for LA-ICPMS are of variable size, between 80 and $350\mu\text{m}$. They are stubby to multifaceted and prismatic, and show a light pink to red color. Under cathodoluminescence they generally present one or two episodes of variably luminescent overgrowths, developed around more or less preserved relict cores. The cores often show oscillatory zoning (Fig. 7). U-Pb dating stresses the extreme complexity of such samples, and the record of several distinct events. The inherited cores are generally concordant, and broadly Grenvillian in age, ranging between ~900 to ~1100Ma (Fig. 8F). Their Th/U ratios ($\gg 0.01$) as well as the REE pattern, with marked positive Ce and negative Eu anomalies, strongly confirm the igneous nature of the protolith (Fig. 9C and Table I (electronic appendix in www.geologica-acta.com), (e.g. Hoskin and Schaltegger, 2003). The data obtained on the overgrowth analyses indicate the presence of at least two different events. A first recrystallization event occurred during the Mississippian (324-348Ma). The second overgrowth developed during the Triassic - Latest Permian

(218-253Ma), generating thin oscillatory bands under CL (Fig. 7). In all the overgrowth episodes the corresponding

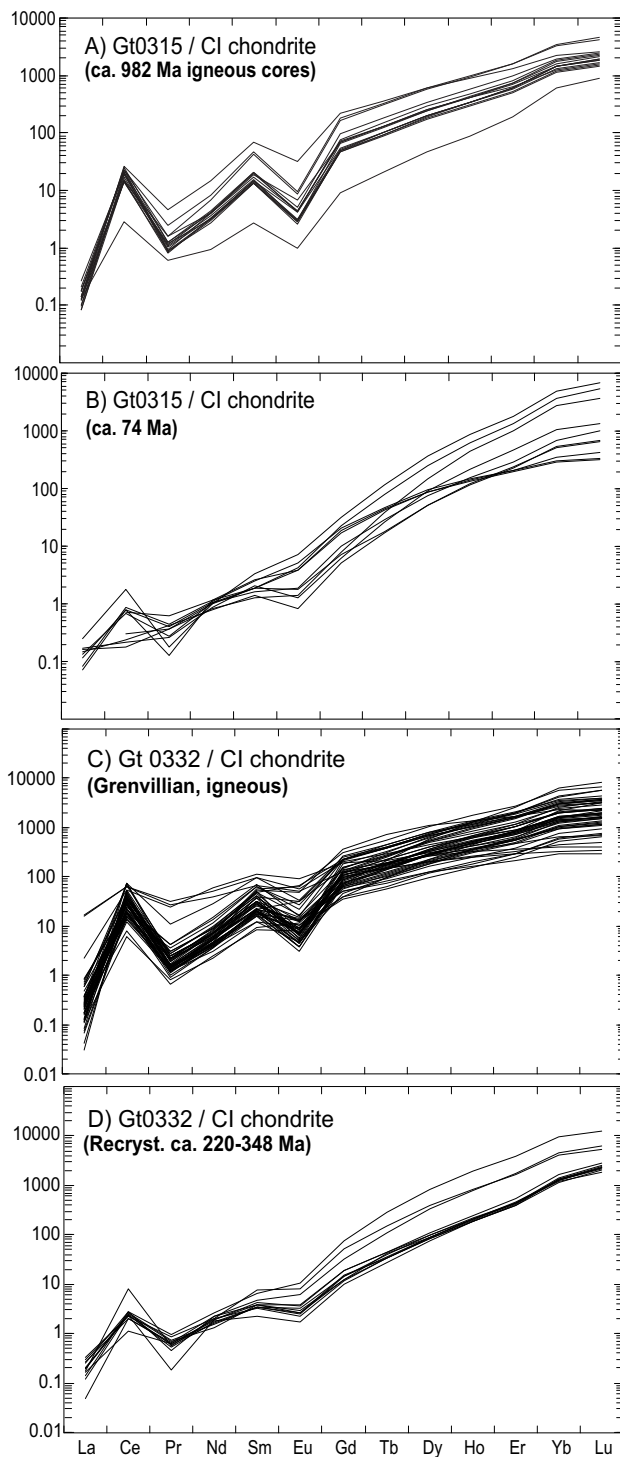


FIGURE 9 | Zircon REE elements patterns normalized to the chondrite values of McDonough and Sun (1995), obtained by LA-ICPMS on the same spot analyzed for U-Pb geochronology. A) ~982Ma igneous cores, pegmatite sample Gt0315; B) ~74Ma metamorphic zircons, pegmatite sample Gt0315; C) Grenvillian igneous zircons, sample Gt0332; D) 220-348Ma recrystallized metamorphic rims, sample Gt0332.

Th/U ratios are lower than igneous cores, and their REE patterns indicate an overall depletion in LREE (La to Eu) with respect to HREE (Gd to Lu), indicating the possible interaction with high-grade metamorphic fluids (Fig. 9D).

Gt 0309

This sample is a pelitic metasedimentary rock made up of garnet, kyanite, phengite, quartz, biotite, plagioclase and abundant rutile. It crops out in the southern Sierra de Chuacús (Fig. 2), north of the unpaved road which connects Pachalúm with Caquíl. Separated zircons range between 40 and 320 μ m in size. They show different colors, varying between dark to pale red, and colorless to amber. They also show a different degree of roundness. One hundred crystals were analyzed, one analysis each, and 6 of them were discarded due to >10% discordance. Of the remaining grains, several yield Proterozoic frequency maximums at 915, 941, 961, and 1106Ma (Fig. 8G). Some younger grains are also concordant (e.g. between ~690 to ~702Ma), however the vast majority are slightly discordant. The lower intercept of such discordant grains is 220 \pm 5Ma. We interpret this lower intercept as indicative of the thermal event responsible for metamorphism and deformation of the sedimentary protholith, as well as for Pb loss in the variably discordant zircons.

Gt 0341

This orthogneiss crops out along the paved road which connects El Rancho with Cobán, near the turn off to Salamá (Fig. 2). It is made up of quartz, plagioclase, k-feldspar augens mantled by biotite and minor white mica. Pyrite is quite abundant as an opaque mineral. The concentrated zircons are colorless to pale colored, with yellow to orange tonalities. A total of 50 analyses were performed on 50 crystals and only one was discarded because of having >10% discordance. The remaining crystals define a principal cluster around 440Ma (concordia of Fig. 8H), with a weighted $^{206}\text{Pb}/^{238}\text{U}$ mean of 449 \pm 1/-3Ma, which is interpreted as the crystallization age of this sample (Fig. 8H inset). Few zircons are inherited from the source, and only two yield younger discordant ages, probably as a result of Pb loss.

Gt 03113

Sample Gt03113 is a granitic dike that cuts across high-grade metamorphic rocks (garnet+amphibole) in the northern Chuacús Complex (Fig. 2). Few zircons were concentrated from this sample. They are dark red, short prismatic, and in general show abundant cracks and inclusions. Only 30 analyses were performed on 30 selected crystals, 3 of which were discarded because of >10% discordance. The remaining analyses define a

well constrained upper intercept at 1096 ± 17 Ma, several variably discordant grains younger than 1100 Ma and also a well constrained lower intercept at 466 ± 27 Ma, underlined by two analyses of crystal rims (Fig. 8I). Two interpretations are possible: either the upper intercept defines the crystallization age of the pegmatite, and the lower is coincident with a tectonothermal disturbance, or the upper intercept represents the inheritance from the protolith, and the lower intercept is the true crystallization age. Several features suggest that the last interpretation is correct: 1) all the analyzed grains have a high Th/U ratio, indicative of igneous crystallization; 2) the pegmatite is undeformed and does not show any metamorphism; and 3) its outcrop is in close vicinity with sample Gt0341, which has a similar age, and with the Rabinal granite suite, which also yielded Ordovician crystallization ages (e.g. Ortega-Obregón et al., 2008).

Gt 0405

This is a migmatitic gneiss which crops out in the southern Sierra de Las Minas, which has been considered as the eastern continuation of the Chuacús Metamorphic Complex. It is exposed along the Pasabien River, 5 km northwest of Río Hondo. The gneiss is made up of up to 20 cm thick leucocratic bands (quartz - plagioclase feldspar - biotite \pm muscovite) alternated with garnet - amphibole - biotite plagioclase horizons. Abundant zircons were separated from this sample. They range from 40 to $350 \mu\text{m}$ in size, and are colorless to intensely red-amber colored. A total of 55 LA-ICPMS analyses were performed on 55 selected grains, 14 of which were rejected because of more than 10% discordance (Table I, electronic appendix in www.geologica-acta.com). The remaining analyses define a bimodal distribution in the concordia diagram of Fig. 8L, with two principal peaks at $450 \pm 2/-4$ Ma and 243 ± 3 Ma. Whereas the older age group corresponding to igneous cores (see cathodoluminescence images of Fig. 7) is readily interpreted as indicating an Ordovician magmatic event, the younger age group is more complex, being characterized by a series of 17 analyses straddling concordia between 255 and 220 Ma. Moreover, those analyses are probably a mix of different zircons: igneous zircons characterized by Th/U ratios >0.1 and a mean $^{206}\text{Pb}/^{238}\text{U}$ age of 249.5 ± 2 Ma and dark red, high-U, metamict zircons with a Th/U ratio <0.02 and $^{206}\text{Pb}/^{238}\text{U}$ ages comprised between 243 and 220 Ma. These last ages reflect a high-grade event maybe coincident with migmatization, which is responsible for the variable Pb loss in some of the analyzed zircons.

DISCUSSION

The new data presented in this paper on the Chuacús Metamorphic Complex allow a better definition of the

tectonothermal events that affected it, as well as its possible paleogeographic settings. Three main events are discussed: 1) Ordovician magmatism, 2) Triassic magmatism and metamorphism and its tectonic implications, and 3) Late Cretaceous magmatism and metamorphism. The autochthonous vs. allochthonous nature of the Chuacús Metamorphic Complex with respect to the Maya Block will be discussed at the end of this chapter.

Ordovician magmatism

Although Grenvillian and other Mesoproterozoic inherited ages are present in most rocks, they are in general overprinted, or reworked by Ordovician magmatism at the southern edge of the Maya Block, as previously reported for the Rabinal-Salamá area of central Guatemala by Ortega-Obregón et al. (2008) who dated S-type granites between 455–462 Ma. Ordovician bodies are also present nearby the Chiapas Massif of SE México, where Pompa-Mera et al. (2008) dated an S-type granite at 482 ± 3 Ma and an amphibolite at 456 ± 14 Ma (U-Pb, single-grain zircons), as well as in the Altos Cuchumatanes, where Solari et al. (2010a) reported an undeformed granite whose crystallization age was dated at $461 \pm 6/-3$ Ma. In Belize the magmatism is mainly Silurian, dated at ~ 410 Ma (Martens et al., 2010). In the northern Sierra de Chuacús Ordovician magmatism is represented by orthogneisses and pegmatitic dikes (samples Gt0341 and Gt03113, respectively), which intrude greenschist facies metasediments belonging either to the San Gabriel unit or the Chuacús Metamorphic Complex (cf. Solari et al., 2009). Bimodal magmatism is widespread in the Acatlán Complex of southern México and associated to rifting of the Rheic Ocean (e.g. Miller et al., 2007; Ortega-Obregón et al., 2009). In northwestern México the Río Fuerte Formation (Late Ordovician) contains magmatic zircons of both Ordovician and Cambrian age (Vega-Granillo et al., 2008). Ordovician magmatism has been documented as well all around the Rheic Ocean realm, generally represented by continental, within-plate tholeiites and as felsic magmas originated by crustal melting of Neoproterozoic and Mesoproterozoic basement (Murphy et al., 2009, 2010, and references therein). Thus, the presence of some Ordovician rocks in the northernmost Chuacús Metamorphic Complex (sample Gt0341), as well as in the nearby Sierra de Las Minas (sample Gt0405) cannot be used as conclusive evidence to establish its autochthony relative to the Maya Block *sensu stricto* (cf. Ortega-Gutiérrez et al., 2007; Ortega-Obregón et al., 2008).

Upper Triassic magmatism and metamorphism

Ordovician magmatism was not recognized in the higher grade central and southern Sierra de Chuacús, where the main magmatic event is Upper Triassic (218–

226Ma, samples Gt0319, Gt0320, Gt0321) and has an arc-like geochemical signature. This magmatic event is also coincident with the tectonothermal event responsible for migmatization in El Chol (sample Gt0332) and Sierra de Las Minas (sample Gt0405). Because remnants of eclogite-facies mafic rocks are present within these deformed migmatites, and are also recognized as deformed lenses nearby the felsic orthogneisses (sample Gt0321, 218Ma) of Agua Caliente River (cf. Ortega-Gutiérrez et al., 2004), we interpret such Upper Triassic ages to be younger than the eclogite-facies metamorphism of the Chuacús Metamorphic Complex. It is interesting to note that none of the S-type Ordovician granites cropping out north of the Sierra de Chuacús (Rabinal-Salamá area), along the Baja Verapaz shear zone of Ortega-Obregón et al. (2008) contain xenoliths of eclogite-facies mafic rocks. This suggests that the eclogite facies metamorphism is constrained between the age of emplacement of the Ordovician granites and the Triassic arc magmas.

Late Cretaceous magmatism and metamorphism

Late Cretaceous metamorphic overprint is present in many of the Chuacús Metamorphic Complex rocks (e.g., Ortega-Gutiérrez et al., 2004, and references therein; Martens et al., 2007; Raschbacher et al., 2009). Some pegmatites were also generated by high-grade metamorphism during the Late Cretaceous, and sample Gt0315 is a clear example of this. Late Cretaceous metamorphism was also reported in the San Gabriel unit (Ortega-Obregón et al., 2008), along the Baja Verapaz Shear Zone, during overthrusting of the Chuacús Metamorphic Complex on top of the southern Maya Block edge. High-pressure metamorphism is also evidenced in several Ar-Ar ages calculated in phengites belonging to the Motagua mélange, which yielded 77-65Ma (Harlow et al., 2004) and were interpreted as the result of the Late Cretaceous collision of the Nicaragua Rise with the southern edge of the Maya Block. Tectonic convergence between the northern Caribbean and the southern North American plates caused subduction of the Chuacús Metamorphic Complex and HP-HT conditions followed by its exhumation, as well as by the northward obduction of El Tambor ophiolites and closure of the Proto-Caribbean Ocean along the paleo Motagua Suture Zone (e.g. Brueckner et al., 2009; see also García-Casco et al., 2008 for correlation with the Antilles).

Allochthony vs. autochthony of the Chuacús Metamorphic Complex with respect to the Maya Block

The Chuacús Metamorphic Complex constitutes the highest-grade metamorphic complex of all Central America. Major faults bound the complex to the north (Baja Verapaz) and south (Motagua fault system), and

no similar rocks occur beyond those faults. Although the Chuacús Metamorphic Complex occurs next to typical Maya cover and basement units of pre-Jurassic age, its youngest cover is of Neogene age, and thus, in principle, it should be considered a suspect tectonostratigraphic unit presently located between the Maya and Chortís blocks. At least three points must be considered to further evaluate the allochthonous versus autochthonous nature of the Chuacús Metamorphic Complex with respect to the Maya Block.

The Maya Block north of the Baja Verapaz Shear Zone and, furthermore, north of the Polochic fault, does not record many of the tectonothermal events that have been recognized in the Chuacús Metamorphic Complex. These include HP metamorphism (e.g. Ortega-Gutiérrez et al., 2004; Ratschbacher et al., 2009, and this paper), Triassic magmatism and deformation/migmatization (Ratschbacher et al., 2009, and this paper), and widespread presence of Late Cretaceous, amphibolite facies reworking and magmatism (McBirney, 1963; Ortega-Gutiérrez et al., 2004; Martens et al., 2007; Martens, 2009; Ratschbacher et al., 2009, and this paper).

No other basement outcrop within the Maya Block (Guichicovi, Chiapas Massif, Maya Mountains) shares geological similarities with the Chuacús Metamorphic Complex. For instance the Guichicovi Complex is composed of Grenvillian (1000-1300Ma, Weber and Kohler, 1999) granulites, only affected by Permian intrusions; whereas the Chiapas Massif is made up of high temperature gneisses, intruded by Late Permian granites metamorphosed at 250-254Ma (Weber et al., 2005), and later covered by Jurassic red beds (Blair, 1988). The Maya Mountains are constituted by Silurian plutons, Devonian (and older) strata, covered by Early Devonian volcanics and sedimentary rocks of late Paleozoic age (Martens et al., 2010). None of the aforementioned units contain high pressure rocks. Furthermore the Chuacús Metamorphic Complex is characterized by the absence of Devonian and Permian ages, as well as by the absence of any sedimentary cover older than the Cenozoic.

The voluminous Late Permian granitic magmatism, recently dated by Schaaf et al. (2002) and Weber et al. (2005) in the Chiapas Massif, as well as by Solari et al. (2010a) in the Altos Cuchumatanes, is not represented by detrital zircon ages in any of the more than 500 zircons dated in this work. Furthermore, volcanic units in the Jurassic Todos Santos Formation, cropping out in Chiapas state, and farther to west, in the Cuicateco terrane of southern México, were recently dated by Godínez-Urban et al. (2011) and Pérez-Gutiérrez et al. (2009), respectively, and yet none of the dated samples show the presence of Middle-Upper Triassic zircons,

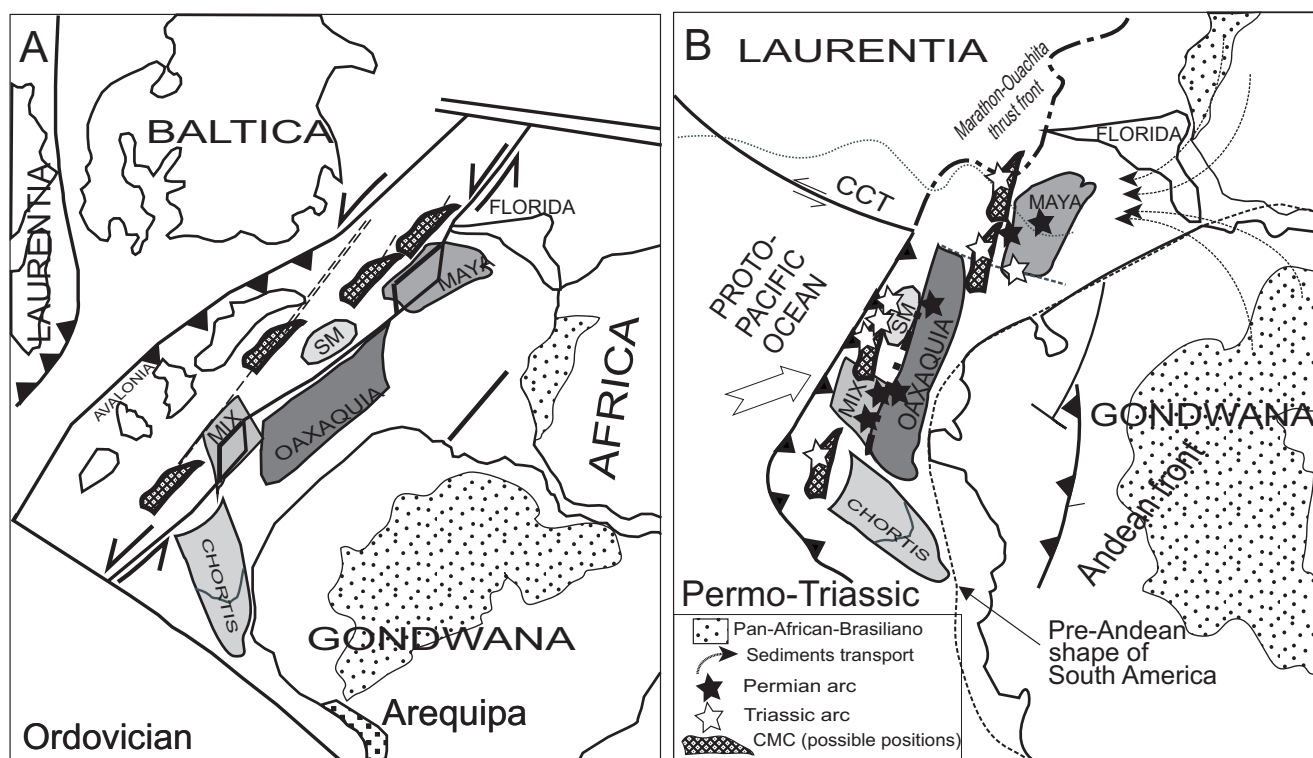


FIGURE 10 | Paleogeographic models for the Ordovician A) and the Permo-Triassic B) Modified from Dickinson and Lawton (2001), Ortega-Obregón et al. (2008), Nance and Linnemann (2008), Weber et al. (2008), and Solari et al. (in press). MIX: Mixteco terrane; SM: Sierra Madre terrane; CCT: California-Coahuila Transform.

such as those dated (213–226Ma) in this work. Although these features may not be considered as conclusive and further studies will be needed, they strongly support the allochthony of the Chuacús Metamorphic Complex with respect to the Maya Block.

On the other hand, correlation of the Chuacús Metamorphic Complex with the Chortís Block to the south of the major Guatemalan faults may be also discarded because the Chortís Block lies in on a different plate (the Caribbean plate, which is considered a far traveled unit originated in the Pacific). Moreover basement units in this block that are most proximal to the Chuacús Metamorphic Complex radically differ in fundamental geologic aspects, such as protoliths and contrasting metamorphic conditions (HP/HT in the Chuacús Metamorphic Complex, versus HT/LP in Las Ovejas Complex, or very low grade metamorphism of the pre-Mesozoic San Diego phyllite, Solari et al., 2009; Ratschbacher et al., 2009, and references therein).

Paleozoic-Triassic paleogeographic reconstructions

Paleogeographic reconstructions for the Late Ordovician–Early Silurian and Permo-Carboniferous suggest that the present southern edge of the Maya Block was located adjacent to northeastern México (e.g.,

Dickinson and Lawton, 2001; Keppie et al., 2006, 2008a; Weber et al., 2008) (Fig. 10A). These interpretations also agree with the Permian paleomagnetic data of Steiner (2005).

In these models, Paleozoic rocks of the Rabinal-Salamá area, just north of our studied area, would have lain close to NE México, north of Oaxaquia and near the Mixteco terrane, where megacrystic granites intruded a passive-margin sedimentary sequence older than ~460Ma (e.g., Nance et al., 2006). Those terranes would have been located close to northwestern Gondwana until the closure of the Rheic Ocean accompanying the amalgamation of Pangea in the late Carboniferous-Permian (e.g., Nance and Linnemann, 2008). The S-type granites cropping out to the north of the studied area could have been produced by partial crustal melting during the Cambro-Ordovician rifting episode coincident with the initial stages of the opening of the Rheic Ocean and the separation of Avalonia from Gondwana (e.g., Murphy et al., 2004).

This suggestion is in agreement with the zipper tectonics model invoked by Martens et al. (2010), which would explain the Silurian-Devonian shifting of magmatism to the Maya Mountains of Belize during subsequent extension related to the opening of the Rheic Ocean. From the Early Pennsylvanian to Permian, the oblique

convergence between Laurentia and Gondwana apparently caused the development of an E-dipping subduction zone with the onset of a new phase of magmatism in the Altos Cuchumatanes (312–317Ma, Solari et al., 2009), and thereafter in northwestern and eastern México (e.g. Torres et al., 1999), the Mixteco terrane (~288Ma, Yañez et al., 1991), the Oaxacan Complex (~275Ma, Solari et al., 2001), and in the Chiapas Massif (~272Ma, Weber et al., 2007).

In this scenario, the Chuacús Metamorphic Complex protoliths could have been deposited in a marginal basin receiving minor input of Ordovician zircons during the Silurian, as demonstrated by their presence in one metasedimentary sample dated by Solari et al. (2009) in the northernmost Sierra de Chuacús (Fig. 10A).

The proposed time constraint on HP metamorphism in the Chuacús Metamorphic Complex (Mid Ordovician to Upper Triassic) is roughly coincident with HP metamorphism reported in the Acatlán Complex of southern México, which was dated as Devonian and was exhumed during the Mississippian (e.g., Keppie et al., 2008). In this view, the Chuacús Metamorphic Complex would have been subducted during the Devonian and metamorphosed at high-grade conditions during the closure of the Rheic Ocean in the late Paleozoic. Exhumation would have taken place causing decompression melting, magmatism and migmatization in the Late Triassic (218–226Ma, Fig. 10B). The Triassic ages in the Chuacús Metamorphic Complex are only slightly younger than those of the Permo-Triassic arc of Chiapas (e.g. Weber et al., 2007), and this can be explained by an easternmost position for the Chuacús Metamorphic Complex in the evolving E-directed, proto-Pacific magmatic arc active since the Permo-Triassic (Fig. 10B).

CONCLUDING REMARKS

The discovery of high-grade metamorphic conditions in the Chuacús Metamorphic Complex of central Guatemala (Ortega-Gutiérrez et al., 2004) has awoken interest and appreciation as a key element for the geologic evolution of the Caribbean region. This contribution presents new geochronologic and geochemical data of some of the most distinctive lithologies of the complex, and provide important constraints to understand this formerly neglected piece of the Caribbean-North America tectonic realm. Indeed, the new results appear to be inconsistent with previous tectonic interpretations that considered the Chuacús Metamorphic Complex to be related to the Chortís-Maya collision, and instead support an independent evolution and an allochthonous provenance. Nevertheless we recognize that further research must be dedicated to address some other important questions that are still poorly understood, or that

remained largely unexplored: Among those we can suggest the following:

The existence of ultrahigh pressure metamorphism (UHP) once suggested by (Ortega-Gutiérrez et al., 2004) for the Chuacús Metamorphic Complex has yet to be confirmed. Nevertheless, intriguing minerals and peculiar textural features, coupled with thermobarometric calculations that approach such conditions, keep this possibility open for future petrologic work in the eclogitic domains that resisted retrogression inside the high-pressure gneisses.

It is still uncertain if the Chuacús Metamorphic Complex experienced only one eclogite-facies metamorphic event. Although there is clear evidence for a mid-late Paleozoic tectonothermal event, a pervasive Late Triassic magmatism and metamorphism as well as an intense Cretaceous tectonothermal overprint throughout most of the sequence; further and more detailed studies are needed to elucidate the kinematic history of its protracted tectonic evolution. In particular, the structural and kinematic history of the Chuacús Metamorphic Complex has remained almost unattended, despite its basic importance.

Finally, the lithotectonic nature of Chuacús Metamorphic Complex protoliths, as preliminary described in this work, should in the near future be an attractive target to unravel what is, perhaps, the most important question that has not been satisfactorily answered: Where does the Chuacús Metamorphic Complex come from, and which were its connections to the now adjacent Maya and Chortís blocks?

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ELECTRONIC APPENDIX

TABLE 1 | Trace and REE concentrations of the selected zircons of Figure 9

Sample number	P	Ti	Y	Nb	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Hf	Th	U	Log(Ti)	T°C
G10315, pegmatite																						
14°58'15.49"																						
90°43'29.88"																						
G10315 002	6	bdl	250	1.44	0.03	0.49	0.01	0.52	0.38	0.22	3.78	1.6	21	7	31	47	8	15449	0.07	54	bdl	
G10315 003	4	13	268	1.56	0.02	0.54	0.04	0.51	0.37	0.28	4.10	1.7	23	8	33	49	8	15998	0.06	57	1.103199066	762
G10315 004	70	bdl	1462	13.6	0.06	1.08	0.02	0.45	0.48	0.41	6.43	4.2	88	46	290	778	171	18943	0.14	945	bdl	
G10315 005	6	2	245	1.57	bdl	0.44	0.06	0.52	0.27	0.21	3.42	1.5	21	7	32	57	10	16040	0.05	52	0.396018219	632
G10315 007	bdl	6	423	2.93	0.04	0.13	0.02	0.40	0.30	0.07	1.21	0.9	23	12	74	167	33	18494	0.02	168	0.763894393	695
G10315 008	93	2	1078	4.53	0.02	0.49	0.04	0.49	0.27	0.24	4.43	2.8	61	33	213	594	131	18351	0.11	688	0.21118804	603
G10315 009	4	bdl	235	1.32	0.03	0.14	0.04	0.36	0.21	0.05	1.00	0.6	13	7	38	82	16	17691	0.01	33	bdl	
G10315 010	73	18	562	4.12	0.02	0.16	0.10	1.48	2.20	0.23	10.8	3.7	49	19	94	211	42	13183	71	189	1.243679001	793
G10315 011	47	25	502	4.52	0.03	0.95	0.07	1.48	2.11	0.16	10.0	3.4	45	17	83	189	37	13579	13	121	1.389978264	827
G10315 013	bdl	3	829	7.94	0.04	0.11	0.03	0.39	0.19	0.08	1.57	1.4	36	24	163	431	89	20941	0.01	82	0.521366919	653
G10315 014	80	19	690	4.16	0.02	0.18	0.15	1.81	2.48	0.25	13.9	4.6	61	23	116	259	51	13735	29	88	1.271995867	799
G10315 015	119	10	1600	6.39	0.04	13.9	0.23	3.71	6.98	0.53	36	11.9	151	54	259	538	103	13842	60	100	0.984846748	738
G10315 016	114	10	798	8.00	0.04	12.8	0.12	2.00	3.01	0.24	15.3	5.4	72	27	133	295	58	13892	36	76	1.010989262	743
G10315 017	13	3	275	1.74	0.03	0.41	0.03	0.47	0.24	0.11	1.98	1.1	19	8	46	111	24	16572	0.02	59	0.454141414	641
G10315 018	62	23	565	5.80	0.06	12.5	0.08	1.24	2.00	0.16	9.50	3.3	48	18	98	232	47	15266	257	1009	1.360248977	820
G10315 019	107	14	732	2.56	7.98	3.0	1.87	7.77	3.65	0.61	16.8	5.4	69	25	119	229	43	11723	86	191	1.151753091	773
G10315 020	67	7	657	3.78	0.03	0.78	0.10	1.91	2.86	0.25	14.2	4.6	62	23	112	235	44	13904	3.05	30	0.843004284	710
G10315 021	18	2	215	1.08	bdl	0.18	0.03	0.46	0.28	0.10	1.42	0.7	12	6	37	85	17	17363	0.02	45	0.339636339	623
G10315 022	93	20	462	4.14	0.04	8.61	0.08	1.40	1.95	0.16	9.18	3.2	43	16	79	182	35	14119	21	90	1.291892951	804
G10315 023	72	19	575	5.46	0.03	9.68	0.11	1.44	2.19	0.17	10.2	3.7	49	19	96	233	47	14421	151	397	1.286472869	802
G10315 024	123	25	1618	6.39	0.03	13.9	0.15	3.03	6.40	0.48	33	11.0	144	52	259	566	111	14055	392	706	1.405922792	830
G10315 025	48	24	562	5.65	0.02	12.0	0.08	1.37	1.97	0.15	9.58	3.6	48	18	95	228	46	13917	285	901	1.381290362	824
G10315 026	73	bdl	162	3.43	0.03	1.74	0.06	0.44	0.41	0.05	1.77	0.7	11	5	31	96	21	15183	34	390	bdl	
G10315 027	72	24	706	5.75	0.02	11.4	0.10	1.63	2.33	0.16	12.8	4.5	60	23	118	277	55	14155	302	901	1.379728486	824
G10315 028	100	13	966	5.16	0.03	15.3	0.12	1.80	3.04	0.28	19.4	6.5	85	32	157	310	59	13086	199	469	1.114091374	765
G10315 029	127	28	1390	2.22	0.03	15.9	0.43	6.82	10.3	1.79	45	12.9	149	48	213	355	65	10947	115	186	1.440732925	839
G10315 030	99	16	639	2.54	0.05	13.8	0.09	1.80	2.82	0.39	13.9	4.6	61	22	103	202	39	11737	71	169	1.199154779	783
G10319, deformed diorite																						
14°57'56.9"																						
90°32'49"																						
G10319 001	88	20	1651	2.06	0.05	30	0.32	4.05	6.05	2.45	30	10	139	52	271	568	112	9578	221	204	1.295550897	805
G10319 003	58	bdl	724	1.29	0.01	12	0.17	2.02	2.69	1.18	12	4	58	23	124	304	68	10035	184	203	bdl	
G10319 004	32	8.95	649	1.19	0.01	14	0.06	1.02	1.55	0.77	9	3	47	19	108	284	67	12442	165	238	0.951763132	731
G10319 005	64	bdl	896	2.00	0.06	18	0.33	3.47	4.32	2.13	16	5	68	27	147	371	83	8458	282	234	bdl	
G10319 006	173	bdl	2516	5.77	0.19	50	0.72	8.84	11.16	6.19	49	16	206	76	397	924	192	7752	991	549	bdl	
G10319 007	185	10	2039	3.63	0.19	31	0.74	10.39	12.52	6.06	47	14	176	64	322	709	146	6812	316	184	0.982921036	738
G10319 008	64	1.19	456	1.02	0.02	9.38	0.09	1.34	1.31	0.62	5.87	2.16	30	13	78	226	55	8655	79	99	0.077303038	583
G10319 009	48	bdl	190	0.62	0.01	4.06	0.05	0.82	0.63	0.26	2.49	0.85	13	6	35	113	29	8004	15	39	bdl	
G10319 010	148	4.53	1996	4.27	0.13	40	0.64	8.63	11.09	4.89	42	13	171	62	318	710	149	7812	470	280	0.656560103	676
G10319 011	62	3.28	542	1.00	0.03	9.71	0.24	2.65	2.63	1.13	10	3	42	16	88	229	51	7199	72	69	0.515478469	652
G10319 012	69	bdl	491	1.13	0.00	10.4	0.06	1.15	1.50	0.69	8	3	37	15	83	205	45	10755	95	113	bdl	
G10319 014	16	5.65	239	1.08	bdl	6.56	0.03	0.64	0.54	0.22	2	0.98	15	7	43	134	32	11627	31	87	0.752050631	693
G10319 015	19	bdl	663	1.08	0.03	11.7	0.09	1.57	2.35	1.08	12	4	54	21	111	252	54	10398	193	204	bdl	
G10319 016	42	17	817	2.09	0.03	14.8	0.14	1.85	2.24	0.97	12	4	64	25	140	338	73	9560	247	294	1.240889137	792
G10319 017	104	8.82	1053	2.67	0.10	25	0.35	4.36	4.67	2.32	19	6	82	32	171	429	93	8463	344	259	0.945369276	730
G10319 018	34	6.60	526	1.14	0.02	8.97	0.06	1.23	1.38	0.70	7	3	37	16	90	243	61	10957	128	232	0.819354015	706
G10319 019	24	bdl	470	1.07	0.04	9.01	0.05	1.13	1.47	0.76	7	3	35	15	80	206	47	10977	131	169	bdl	
G10319 021	93	bdl	831	2.18	0.08	19.6	0.23	2.33	2.83	1.42	11	4	58	24	141	389	89	8269	250	247	bdl	
G10319 022	30	bdl	41	0.46	0.01	1.54	0.04	0.49	0.24	0.03	0.34	0.12	2.03	1.14	8	36	11	11556	3	43	bdl	
G10319 023	124	7.05	2233	3.61	0.07	59	0.38	5.26	8.38	3.52	44	15	196	72	362	690	133	10077	506	293	0.84829666	711
G10319 024	161	32	2044	3.03	0.04	20	0.40	5.39	7.70	3.13	40	14	182	68	334	651	128	8495	119	103	1.498371018	853

TABLE 1 | Continued

Sample number	P	Ti	Y	Nb	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Hf	Th	U	Log(Ti)	T°C
G0319 025	192	7.65	839	2.13	0.00	11.3	0.08	1.30	2.07	0.90	12	4	65	26	144	322	66	8386	28	44	0.883559191	718
G0319 026	150	8.93	1784	3.77	0.15	23	0.49	5.12	7.14	3.91	35	11	144	54	284	686	147	7112	463	292	0.950930639	731
G0319 027	115	bdl	572	1.53	0.00	11.4	0.05	0.90	1.27	0.48	7	3	41	18	103	267	61	10293	145	228	bdl	
G0319 028	16	5.80	125	1.03	0.03	6	0.05	0.78	0.24	0.10	1.27	0.49	8	4	23	71	18	11740	80	207	0.763214558	695
G0319 029	103	5.66	1616	3.09	0.08	16	0.25	3.58	5.26	2.42	30	10	140	53	265	535	106	8024	86	81	0.752565659	693
G0319 030	187	8.87	2536	9.27	0.02	38	0.28	3.94	6.79	2.55	37	14	200	82	435	907	182	8523	164	170	0.947703735	730
G0319 031	63	bdl	1014	2.01	0.05	16	0.19	2.72	3.30	1.07	19	7	88	33	172	362	74	9949	188	185	bdl	
G0319 033	146	1.42	1077	2.32	0.03	26	0.12	1.72	2.55	1.17	14	6	82	33	186	428	90	9270	166	180	0.153732124	594
G0319 034	87	6.10	899	2.27	0.08	16	0.21	2.43	3.42	1.84	16	5	69	27	145	384	83	7580	225	174	0.785504443	699
G0319 037	99	bdl	1379	1.87	0.04	15	0.23	3.72	5.82	2.40	28	9	123	45	228	448	87	8719	66	63	bdl	
G0319 038	72	bdl	873	1.45	0.07	16	0.29	3.67	3.44	1.70	16	5	69	26	140	350	75	8042	188	175	bdl	
G0319 039	37	3.59	951	1.65	0.02	18	0.20	2.45	3.34	1.85	16	5	69	26	140	350	75	8042	188	175	bdl	
G0319 040	42	3.87	351	0.87	0.06	7.1	0.09	1.31	1.07	0.55	5	1.74	25	10	60	172	41	8855	51	72	0.55558113	658
G0319 041	49	11	675	1.24	0.07	13	0.29	3.30	3.43	1.43	13	4.08	54	20	110	274	59	7633	119	110	0.588209591	664
G0319 042	10	bdl	180	0.70	0.02	4.52	0.05	0.67	0.40	0.18	1.74	0.70	11	5	32	93	24	12253	56	97	1.02155217	745
G0319 043	88	340	956	1.43	0.07	8.23	0.26	2.91	3.82	2.01	19	6	80	30	154	344	74	6633	151	135	bdl	
G0319 044	68	9.48	1212	2.43	0.10	21.17	0.37	4.23	5.24	2.70	22	7	98	37	196	468	100	7581	273	214	2.530947298	1187
G0319 045	100	bdl	1766	2.24	0.05	16.44	0.33	5.41	7.75	2.83	36	12	159	58	289	560	109	7963	91	82	0.976713284	736
G0319 046	120	5.43	1861	3.47	0.13	34	0.55	6.65	8.35	3.83	37	12	155	57	298	682	139	7454	407	294	bdl	
G0319 047	48	bdl	805	1.73	0.07	12.89	0.21	2.73	2.92	1.51	14	5	63	25	131	328	71	9237	217	170	0.734849388	690
G0319 048	68	12	462	1.27	0.03	8.93	0.14	1.64	1.58	0.63	6	2.31	33	14	78	207	47	7612	63	68	bdl	
G0319 049	47	bdl	437	0.94	0.05	8.64	0.07	1.14	1.25	0.63	6	2.13	32	13	78	215	50	9933	86	126	1.069009689	755
G0319 050	172	1.40	2148	2.90	0.07	17.49	0.43	6.03	9.02	3.56	44	15	198	72	358	669	132	7801	124	93	0.147002165	593
G0319 051	203	bdl	1407	3.88	0.03	30	0.17	2.47	3.28	1.68	18	7	103	43	241	591	126	7712	322	292	bdl	
G0319 053	20	bdl	180	0.75	0.05	4.27	0.06	0.76	0.71	0.28	3	0.95	13	5	32	93	23	9607	53	70	bdl	
G0319 054	123	6.16	1707	3.91	0.15	31	0.49	6.35	8.40	4.12	35	11	141	52	271	645	135	7897	497	305	0.789745011	700
G0319 055	36	5.10	1010	1.76	0.08	17	0.30	3.31	3.59	1.69	17	6	78	31	166	395	91	8902	279	248	0.707723998	685
G0319 056	81	bdl	1366	1.68	0.10	13	0.35	3.89	5.38	2.66	26	9	115	43	221	509	108	7901	263	228	bdl	
G0319 057	68	bdl	1276	1.30	0.01	2.2	0.20	3.12	4.49	2.10	23	8	108	42	209	436	89	10451	201	198	bdl	
G0319 058	23	bdl	525	1.09	0.06	11.28	0.06	0.86	1.43	0.79	9	2.96	41	17	90	217	48	11704	110	144	bdl	
G0319 059	134	7.97	2173	4.47	0.10	39	0.56	6.19	8.28	4.46	43	14	184	68	350	762	153	9058	560	410	0.901251809	721
G0319 060	80	10	766	1.42	0.10	12.2	0.37	3.79	3.81	1.95	15	5	62	23	127	311	67	7068	103	86	1.000299154	741
G0320, deformed gabbro																						
14°57'56.9"																						
90°32'49"																						
G0320 001	524	7.70	1756	2.98	0.06	19.27	0.36	4.92	7.50	3.38	39.02	12.09	158	58	271	595	107	7433	349	268	0.886218314	718
G0320 002	578	9.91	1391	2.28	0.14	16.54	0.37	4.41	5.79	2.84	30.80	9.47	121	45	212	468	88	7353	249	204	0.995998407	740
G0320 003	412	10.50	1117	6.45	0.05	23.12	0.14	1.81	1.92	1.12	12.64	4.88	75	34	192	517	103	11243	270	404	1.021189916	745
G0320 006	41	bdl	130	0.58	0.03	4.02	0.07	1.29	0.62	0.12	1.62	0.53	8	4	23	78	19	10373	31	58	bdl	
G0320 008	315	7.87	431	1.46	0.08	10.72	0.13	1.21	1.19	0.63	6.33	2.29	33	13	71	186	37	10434	44	60	0.896249179	720
G0320 009	564	0.64	1639	2.63	0.13	24.07	0.43	6.08	8.41	3.68	36.95	11.59	143	53	248	502	90	7519	256	199	-0.19603029	546
G0320 010	849	7.80	2218	4.46	0.11	35.60	0.57	7.05	10.23	4.47	49.97	15.72	197	71	336	680	124	6935	398	295	0.892123547	720
G0320 011	554	0.86	1924	2.97	bdl	24.32	0.39	6.18	8.62	3.40	43.31	13.33	173	62	295	562	100	9388	167	140	-0.0659347	563
G0320 012	664	1.28	2326	3.96	0.08	28.89	0.46	7.74	9.35	3.81	50.07	16.17	205	76	358	687	121	9447	224	181	0.105606177	587
G0320 013	247	7.39	664	2.66	0.07	11.49	0.10	1.57	1.69	0.73	8.43	3.16	45	20	114	308	66	8907	100	155	0.868745364	715
G0320 014	460	bdl	689	2.58	0.02	11.42	0.12	1.48	1.86	0.73	8.11	3.07	47	21	116	314	66	8476	95	140	bdl	
G0320 016	770	11.66	2478	4.17	0.06	29.20	0.44	6.10	8.63	3.82	47.82	16.43	214	82	382	774	133	8700	378	289	1.066725496	755
G0320 018	158	2.22	402	1.85	0.02	9.51	0.10	1.63	1.30	0.55	6.66	2.21	31	13	68	186	37	9413	41	88	0.346833695	624
G0320 020	967	3.30	1477	7.11	0.00	27.02	0.08	1.67	2.90	1.13	19.37	7.13	108	47	246	622	117	12343	193	281	0.517852767	652
G0320 021	244	5.77	1061	2.04	0.05	13.79	0.22	3.47	4.52	2.13	22.08	6.91	90	35	171	399	74	9554	209	194	0.761329923	695
G0320 022	196	bdl	414	0.88	bdl	5.69	0.15	1.79	1.55	0.51	6.28	1.99	31	13	75	235	49	10468	85	178	bdl	
G0320 023	356	6.06	637	2.92	0.03	10.94	0.13	1.70	1.75	0.81	8.79	3.15	46	20	108	295	59	9583	119	191	0.782565847	699
G0320 025	250	0.51	610	1.14	0.05	9.70	0.15	2.10	2.27	1.10	11.99	3.73	49	20	96	233	45	9216	123	135	-0.293224655	533
G0320 026	480	14.25	654	2.45	0.07	13.08	0.15	1.90	2.17	1.06	10.33	3.55	50	21	111	282	55	9652	130	185	1.153810115	773
G0320 027	378	bdl	852	2.11	0.11	16.75	0.38	4.55	4.89	2.26	20.44	6.40	78	28	131	287	50	7860	192	150	bdl	
G0320 028	818	4.08	1164	2.71	0.07	17.81	0.19	2.58	3.85	1.57	19.84	6.85	93	37	187	432	79	7164	121	131	0.610500573	668
G0320 029	149	10.04	216	0.84	0.07	4.98	0.08	1.32	1.05	0.44	4.54	1.45	19	7	35	82	15	7640	53	66	1.001905299	741

TABLE I | Continued

Sample number	P	Ti	Y	Nb	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Hf	Th	U	Log(Ti)	T°C
GI0320 030	130	bdl	296	0.76	0.03	5.94	0.10	1.10	1.11	0.63	5.54	1.83	25	10	48	113	21	7954	88	99	bdl	
GI0320 032	369	6.90	962	2.18	0.09	15.65	0.31	3.83	4.89	2.23	20.47	6.42	84	30	149	350	64	7961	162	139	0.839156302	709
GI0320 033	327	2.09	676	1.16	0.06	9.26	0.21	2.50	2.85	1.39	12.38	4.22	54	21	110	282	55	8407	103	117	0.320630029	620
GI0320 034	441	4.23	517	2.13	0.07	11.16	0.10	1.31	1.44	0.56	7.15	2.44	38	16	88	246	49	8373	74	160	0.62597851	671
GI0320 035	187	7.00	378	1.32	0.02	6.34	0.11	1.59	1.64	0.65	6.50	2.15	31	12	64	177	35	8130	51	79	0.8444869035	711
GI0320 036	388	6.88	947	2.37	0.02	14.36	0.21	3.02	2.89	1.58	14.34	5.33	73	30	160	413	86	10503	330	382	0.837443497	709
GI0320 037	547	1.79	472	1.22	0.01	8.24	0.10	1.26	1.31	0.74	6.46	2.42	36	15	80	227	47	9303	103	148	0.25242783	609
GI0320 038	1052	6.33	3483	7.64	0.22	57.42	0.82	11.28	15.94	6.87	83.92	25.90	326	116	531	1047	178	8457	534	323	0.801597915	702
GI0320 039	543	7.70	1409	2.73	bdl	24.79	0.18	2.36	3.92	1.86	25.42	8.70	118	47	233	526	97	13112	464	450	0.866220897	718
GI0320 040	443	bdl	1673	3.36	0.04	24.84	0.21	3.26	5.84	2.67	31.32	10.66	145	55	271	604	111	12329	539	484	bdl	
GI0320 041	383	3.29	639	3.20	0.00	16.22	0.09	1.23	1.52	0.72	8.09	2.96	44	20	112	311	63	12069	234	359	0.517786085	652
GI0320 042	430	2.58	1199	3.03	0.09	16.41	0.31	4.11	5.22	2.61	27.54	8.27	107	39	186	421	73	7530	293	216	0.411205829	634
GI0320 043	748	2.08	1260	3.66	0.00	20.40	0.15	2.18	3.44	1.49	17.83	6.89	98	41	213	552	106	10944	328	371	0.318707321	620
GI0320 044	471	5.72	811	1.50	0.04	9.71	0.23	2.84	3.51	1.58	16.81	5.55	72	27	129	295	53	7938	119	122	0.757695036	694
GI0321, Agua Caliente River felsic orthogneiss																						
14°56'07.8"																						
90°30'10.9"																						
GI0321 006	-	5.5	2375	-	0.07	58	-	-	5.5	1.33	-	-	183	-	-	-	175	12389	226	440	0.743273313	692
GI0321 007	-	7.6	1082	-	0.44	37	-	-	3.0	0.91	-	-	82	-	-	-	87	12656	98	182	0.882269836	718
GI0321 009	-	bdl	881	-	0.01	32	-	-	2.4	0.57	-	-	71	-	-	-	67	12170	47	75	bdl	
GI0321 010	-	0.85	1609	-	0.05	49	-	-	2.5	0.56	-	-	95	-	-	-	187	15203	753	955	-0.072605489	562
GI0321 013	-	5.6	2491	-	0.06	61	-	-	7.1	1.79	-	-	202	-	-	-	146	12472	159	197	0.74612956	692
GI0321 016	-	10	2500	-	0.10	59	-	-	7.4	1.84	-	-	203	-	-	-	159	10133	176	168	1.00441451	742
GI0321 019	-	5.4	1800	-	0.11	51	-	-	3.9	1.02	-	-	135	-	-	-	139	12884	189	266	0.729411984	689
GI0321 020	-	1.1	1967	-	0.06	62	-	-	3.8	0.89	-	-	152	-	-	-	155	14318	594	649	0.033785807	577
GI0321 021	-	7.4	4358	-	1.3	113	-	-	12	2.86	-	-	356	-	-	-	269	12918	694	774	0.867115224	715
GI0321 022	-	9.9	1416	-	0.05	57	-	-	3.1	1.0	-	-	112	-	-	-	105	12360	183	221	0.994758068	740
GI0321 027	-	13	2666	-	39	153	-	-	26	8.12	-	-	217	-	-	-	185	10969	172	220	1.116793241	765
GI0321 030	-	7.2	580	-	0.11	21	-	-	1.2	0.36	-	-	45	-	-	-	52	13687	53	92	0.857475871	713
GI0321 033	-	8.0	1392	-	0.54	38	-	-	3.2	0.42	-	-	108	-	-	-	118	15678	307	474	0.900545554	721
GI0321 034	-	4.1	1464	-	0.24	32	-	-	4.3	0.77	-	-	120	-	-	-	96	14191	156	386	0.617962445	669
GI0321 037	-	3.6	1478	-	1.7	61	-	-	5.5	1.65	-	-	119	-	-	-	101	12309	204	206	0.5559389	658
GI0321 044	-	7.1	1816	-	0.05	60	-	-	2.6	0.71	-	-	106	-	-	-	215	15017	986	1181	0.849242039	711
GI0321 045	-	2.7	1734	-	3.7	84	-	-	4.2	1.07	-	-	126	-	-	-	148	11944	522	505	0.429099049	637
GI0321 048	-	8.1	3083	-	0.30	59	-	-	11	3.48	-	-	260	-	-	-	197	11192	206	208	0.910885653	723
GI0321 049	-	4.1	928	-	0.04	38	-	-	2.5	0.70	-	-	74	-	-	-	76	14300	120	189	0.611383906	668
GI0321 051	-	5.8	1334	-	0.03	46	-	-	2.4	0.65	-	-	93	-	-	-	124	13930	378	460	0.761619418	695
GI0321 052	-	11	2001	-	3.0	81	-	-	6.9	2.47	-	-	151	-	-	-	168	9694	221	298	1.038444237	749
GI0321 055	-	6.6	2022	-	0.08	81	-	-	3.9	0.92	-	-	138	-	-	-	189	13765	764	777	0.817890144	705
GI0321 056	-	bdl	2346	-	0.10	69	-	-	3.9	0.69	-	-	176	-	-	-	187	13721	330	499	bdl	
GI0321 057	-	bdl	1594	-	19	95	-	-	5.2	1.05	-	-	110	-	-	-	146	13885	799	725	bdl	
GI0321 058	-	bdl	1720	-	16	92	-	-	11	2.96	-	-	120	-	-	-	165	14579	517	660	bdl	
GI0321 060	-	8.9	1279	-	8.3	63	-	-	4.6	0.97	-	-	102	-	-	-	99	13504	159	229	0.951678187	731
GI0321 061	-	5.5	1180	-	0.14	34	-	-	3.7	0.43	-	-	101	-	-	-	81	14964	105	238	0.741255368	691
GI0321 062	-	13.9	2190	-	7.2	93	-	-	5.8	1.46	-	-	134	-	-	-	210	14194	1619	1339	1.143553907	771
GI0321 064	-	4.7	657	-	12.9	58	-	-	16	6.65	-	-	60	-	-	-	54	15921	55	187	0.670117775	678
GI0332, El Chol migmatite leucosome																						
14°57'53.6"																						
90°29'12.5"																						
GI0332 001	83	bdl	1914	3	0.05	11	0.09	1.6	3.2	1.8	29	11	158	64	318	697	140	13353	20	204	bdl	
GI0332 002	4	22	2081	2	0.11	10	0.21	3.6	7.2	3.6	44	15	182	69	323	657	141	11779	20	195	1.337857028	814
GI0332 004	49	6	573	3	0.05	11	0.14	2.4	3.1	0.29	13	4.1	50	19	93	210	42	15460	77	382	0.775688502	698
GI0332 005	86	9	807	5	0.02	11	0.15	2.1	2.7	0.48	15	5.4	73	28	133	277	50	13974	7.3	133	0.941255496	729
GI0332 006	20	bdl	2297	13	0.03	1.7	0.08	1.0	1.4	0.63	14	7.8	147	72	403	1005	199	17048	5.2	392	bdl	
GI0332 007	bdl	bdl	412	1.4	0.06	3.3	0.06	0.89	0.64	0.21	4	1.6	24	12	72	225	63	16311	1.6	132	bdl	
GI0332 008	116	bdl	1620	4.6	0.20	13	0.24	3.0	4.2	0.32	27	10	136	54	258	499	90	13121	135	460	bdl	

TABLE 1 | Continued

Sample number	P	Ti	Y	Nb	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Hf	Th	U	Log(Ti)	T ^{°C}
GI0332 009	144	13	3548	55	0.06	1.8	0.09	1.2	1.0	0.60	15	10	210	109	624	1536	314	23711	5.1	1465	1.11469047	764
GI0332 011	176	6	1204	3.8	0.14	12	0.29	4.0	6.6	0.50	30	10	119	43	184	326	58	13412	21	107	0.776792954	698
GI0332 012	83	8	2772	8.6	0.04	17	0.16	3.3	8.2	3.5	60	20	255	95	438	896	166	16607	106	562	0.882341149	718
GI0332 013	88	bid	761	4.2	0.08	8.8	0.13	1.9	2.9	0.25	15	5.4	69	26	123	247	45	13655	50	176	bid	
GI0332 014	77	bid	1918	20	0.04	4.6	0.07	1.4	1.5	0.41	13	7.1	126	62	342	822	171	18907	17	462	bid	
GI0332 015	62	15	437	3.7	0.07	24	0.14	2.6	3.1	0.52	13	4.2	53	19	87	167	32	13272	37	67	1.184900596	780
GI0332 016	20	2	461	4.6	0.06	22	0.25	4.5	7.0	1.0	28	6.8	65	17	58	65	10	14261	16	77	0.298330495	616
GI0332 017	80	bid	1390	3.1	0.04	5.0	0.05	0.93	1.1	0.46	10	5.4	95	46	262	667	131	16816	11	369	bid	
GI0332 018	50	bid	358	1.9	0.08	1.4	0.08	0.91	0.5	0.13	2	1.2	22	12	70	199	45	16861	1.6	108	bid	
GI0332 020	21	17	760	3.6	0.06	4.9	0.06	1.2	1.3	0.44	8	3.4	54	24	132	347	77	15350	25	440	1.223226465	788
GI0332 021	65	3.1	756	3.2	0.02	10	0.11	1.8	2.5	0.21	14	5.1	67	27	125	264	50	13618	54	237	0.496733833	648
GI0332 022	9.1	6.5	323	1.9	0.04	0.7	0.06	0.85	0.33	0.10	2	1.0	19	10	64	218	56	18353	1.3	125	0.81329298	705
GI0332 023	42	bid	1488	5.5	bid	6.0	0.07	1.0	1.2	0.5	12	5.9	103	48	261	652	131	16780	24	307	bid	
GI0332 024	bid	bid	413	1.5	0.04	1.3	0.08	0.79	0.54	0.20	3	1.4	25	13	73	222	63	17270	1.7	136	bid	
GI0332 025	125	1.0	1802	3.1	bid	18	0.10	2.2	5.7	1.8	44	16	193	64	270	524	93	18142	44	453	0.016169753	575
GI0332 026	125	8.0	891	4.5	0.02	31	0.15	2.8	4.6	0.4	24	8.0	97	34	146	238	40	14652	25	96	0.903888831	722
GI0332 027	123	1.8	814	3.9	0.03	33	0.23	3.2	4.6	0.4	21	6.6	81	29	132	251	47	14014	77	68	0.265789719	611
GI0332 028	153	11	1609	7.1	0.07	13	0.39	6.7	10	0.9	46	14	168	59	254	451	79	13858	117	361	1.04310827	750
GI0332 029	156	9.1	2157	8.5	0.03	28	0.13	2.6	7.9	2.7	71	25	271	76	266	372	54	19179	244	696	0.959636191	733
GI0332 030	73	18	256	1.5	0.04	8	0.13	2.1	2.6	0.7	10	2.9	31	9.1	33	47	7.3	12003	0.50	7.1	1.253629001	795
GI0332 031	106	6.2	686	2.0	0.06	45	0.23	3.2	4.6	0.8	20	5.5	67	24	106	199	37	13144	115	111	0.789716169	700
GI0332 033	91	5.3	819	3.0	0.07	11	0.12	2.4	4.0	0.7	21	6.8	84	30	129	230	40	13460	24	96	0.728288268	689
GI0332 034	967	39	7108	24	4.6	34	2.2	15	11	4.2	63	30	493	223	1266	3611	696	33329	44	4137	1.59212312	877
GI0332 035	114	14	778	8.8	0.01	30	0.18	3.1	4.7	0.5	22	6.8	79	28	124	219	40	15717	108	290	1.157080809	774
GI0332 036	2929	17	1878	11	3.8	39	2.9	23	14	1.9	50	14	177	67	306	533	95	13870	195	438	1.232525512	790
GI0332 037	56	bid	302	1.2	0.09	21	0.20	2.8	3.2	0.7	10	2.7	31	11	48	96	18	11363	17	18	bid	
GI0332 038	4754	5.0	635	1.2	0.53	33	2.3	2.8	17	5.0	38	7.8	74	23	93	154	27	10533	47	38	0.694706402	683
GI0332 039	95	2.5	537	6.8	0.03	22	0.23	3.3	5.5	0.3	23	6.6	68	20	75	103	17	14927	19	135	0.39315191	631
GI0332 040	145	11	1658	13	0.08	22	0.20	3.9	6.1	0.4	36	12	154	59	273	494	88	14381	193	406	1.052586702	752
GI0332 041	50	151	374	5.4	0.15	24	0.20	3.4	5.5	0.7	22	5.4	49	14	52	72	12	14819	13	132	2.178634801	1053
GI0332 042	33	18	366	5.4	0.05	18	0.19	3.3	4.7	0.4	20	5.3	51	14	46	52	8	15096	16	168	1.261846578	797
GI0332 043	194	11	2081	5.2	0.20	15	0.39	6.2	9.4	1.2	51	16	202	74	327	563	99	12659	114	218	1.058710446	753
GI0332 044	117	5.3	972	9.5	0.07	16	0.11	1.7	2.7	0.2	17	6.4	85	34	165	332	61	15647	106	370	0.726617297	689
GI0332 045	92	22	976	5.8	0.09	12	0.12	1.9	3.0	0.4	19	6.6	88	35	164	324	59	14811	69	247	1.338619267	814
GI0332 046	153	12	1592	3.5	0.06	9	0.33	5.4	9.4	1.5	46	14	165	58	251	428	77	12165	38	72	1.095834104	761
GI0332 047	192	64	1089	11	0.09	18	0.14	2.0	3.1	0.2	17	6.3	89	37	187	414	81	15475	247	698	1.804898427	935
GI0332 048	130	3.2	1474	8.8	0.03	19	0.14	2.4	4.4	0.3	26	9.2	128	51	245	487	89	15351	195	464	0.506379204	650
GI0332 049	87	22	239	1.3	0.04	20	0.15	1.8	1.8	0.5	7	2.1	23	8.3	39	83	16	12113	5.2	13	1.334855881	814
GI0332 051	74	6	454	2.7	0.05	27	0.14	2.1	2.9	0.6	13	3.7	45	17	75	152	28	12896	37	63	0.749366097	693
GI0332 052	129	bid	956	2.1	0.08	44	1.0	13	14	3.2	41	11	109	35	145	259	45	11846	95	80	bid	
GI0332 053	73	19	379	1.3	0.06	21	0.20	2.8	3.2	0.82	12	3.4	40	14	62	124	23	10248	10	10	1.269906469	799
GI0332 055	109	7.0	779	8.4	0.04	14	0.09	1.7	2.4	0.17	13	5.1	68	27	132	271	49	15701	80	282	0.845070784	711
GI0332 056	147	7.0	708	5.3	0.06	12	0.12	1.8	2.9	0.26	16	5.5	72	26	119	209	38	14217	17	65	0.845419222	711
GI0332 057	196	23	1996	30	0.14	40	0.26	4.3	7.5	0.37	42	14	188	71	324	603	106	14402	315	655	1.36165927	820
GI0332 058	101	7	1303	11	0.17	25	0.29	4.1	6.2	0.36	30	10	127	47	211	386	69	14869	182	467	0.86014947	713
GI0332 059	88	19	715	8.6	0.06	19	0.13	2.4	3.5	0.25	18	5.8	72	26	116	218	39	15180	62	250	1.284174333	802
GI0332 060	93	7.7	1041	6.2	0.04	12	0.15	2.7	4.2	0.47	24	7.9	101	37	171	316	56	14114	54	101	0.885497713	718
GI0332 061	bid	bid	1381	1.6	0.01	1.2	0.06	1.1	0.69	0.34	6.5	4.1	85	44	280	741	156	19846	2.9	305	bid	
GI0332 062	41	19	365	1.3	0.07	1.5	0.05	0.79	0.52	0.18	2.9	1.3	22	11	69	202	54	17636	2.0	174	1.280145715	801
GI0332 063	64	bid	375	1.6	0.05	1.4	0.06	0.74	0.52	0.14	2.9	1.5	24	11	74	210	56	18693	2.3	171	bid	
GI0332 064	31	bid	326	1.2	0.05	1.6	0.07	0.71	0.53	0.15	3.1	1.3	21	10	63	187	51	18171	2.6	264	bid	
GI0332 065	54	171	336	1.4	0.03	1.7	0.04	1.0	0.56	0.15	2.5	1.2	21	10	68	211	58	19682	1.9	279	2.232300797	1072
GI0332 066	bid	bid	437	1.3	0.05	1.5	0.02	0.88	0.58	0.22	3.7	1.6	27	14	86	263	71	19842	1.6	181	bid	
GI0332 067	bid	bid	362	1.1	0.03	1.5	0.06	0.63	0.57	0.17	2.5	1.5	22	11	72	208	57	17884	1.8	141	bid	

GI0405 migmatite Rio Pasabien

15°02'21.4"
89°41'02.4"

TABLE I | Continued

Sample number	P	Ti	Y	Nb	La	Ce	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Yb	Lu	Hf	Th	U	Log(Ti)	T°C	
G10405 002	-	5.99	1026	6.4	0.02	0.02	16	0.07	1.5	3.4	0.78	19	6.8	87	34	157	315	61	12295	160	460	0.777501147	698
G10405 003	-	10	1032	1.5	0.00	3.2	0.09	1.6	3.4	0.16	21	7.1	93	35	156	280	54	12666	38	112	1.006260193	742	
G10405 004	-	13	870	1.5	0.02	9.4	0.12	1.7	3.3	0.50	17	5.7	74	29	135	275	56	11596	98	171	1.116968356	765	
G10405 005	-	16	781	1.8	0.03	5.9	0.08	1.4	2.3	0.18	13	4.9	69	26	126	245	49	11986	78	201	1.207314235	785	
G10405 006	-	21	1236	1.6	0.05	3.9	0.18	2.7	5.3	0.31	26	8.9	115	43	190	354	68	10483	78	193	1.325347534	811	
G10405 007	-	14	1095	4.3	0.26	14	0.18	1.9	3.1	0.15	19	7.0	94	37	177	348	67	13022	204	531	1.139556347	770	
G10405 008	-	bdl	625	1.7	0.01	9.4	0.04	1.0	1.6	0.15	11	3.9	54	21	102	210	41	11412	56	150	bdl		
G10405 011	-	bdl	812	2.4	0.03	9.3	0.08	1.3	2.2	0.24	14	5.1	70	28	129	258	50	11798	80	222	bdl		
G10405 012	-	5.86	1260	5.7	0.34	15	0.34	4.0	5.0	0.27	24	8.6	113	43	203	395	76	13094	224	637	0.767977955	696	
G10405 014	-	12.6	1507	3.1	0.34	18	0.41	4.7	7.6	1.2	38	12	141	50	225	417	79	11634	327	674	1.100910944	762	
G10405 015	-	bdl	1890	9.3	0.70	14	0.38	3.4	5.0	0.57	29	12	165	64	310	676	132	13640	386	2792	bdl		
G10405 017	-	7.98	1277	4.7	0.04	24	0.18	3.3	6.1	0.98	32	9.6	120	44	193	369	71	11760	334	650	0.901998803	722	
G10405 018	-	63	1091	3.9	0.05	13	0.12	1.9	3.8	0.25	21	7.5	100	38	172	333	63	12019	115	250	1.802549922	934	
G10405 019	-	12	2435	4.0	0.61	19.4	0.24	1.20	3.8	3.3	89	22	248	84	350	611	114	12264	585	870	1.092074453	760	
G10405 020	-	1.11	1090	5.1	0.03	16	0.10	1.7	3.6	0.60	21	7.2	93	36	174	375	76	12309	221	794	0.043410358	578	
G10405 021	-	bdl	996	3.7	0.01	15	0.11	1.5	3.1	0.15	19	6.5	88	34	158	312	60	11693	147	359	bdl		
G10405 023	-	12	1067	2.1	5.7	24	1.8	10.0	5.6	0.51	24	7.6	98	37	166	319	61	10494	86	174	1.066278244	755	
G10405 024	-	3.66	727	11	0.14	2.8	0.08	0.82	0.95	0.12	7.1	3.0	51	23	133	386	83	14260	41	1681	0.563937352	660	
G10405 025	-	bdl	673	2.2	0.03	4.2	0.07	1.05	1.2	0.22	10	3.6	53	22	112	278	59	13447	165	492	bdl		
G10405 026	-	bdl	618	3.3	0.02	4.4	0.07	1.05	1.2	0.22	9.4	3.2	49	21	108	277	61	12576	56	608	bdl		
G10405 028	-	5.34	1858	2.5	0.02	12	0.30	5.00	9.5	1.4	49	16	190	69	298	524	99	10135	207	437	0.727853861	689	
G10405 029	-	2.29	1198	5.1	2.87	11	1.01	5.41	2.7	0.15	14	5.9	91	39	198	476	96	15684	130	1802	0.359597789	626	
G10405 030	-	10	1610	2.7	1.91	25	0.89	7.95	10	1.6	44	14	153	54	235	412	79	11647	344	481	1.017094093	744	
G10405 031	-	8.22	1108	2.0	6.87	27	2.75	15.41	7.3	0.63	26	8.3	103	39	174	326	63	9735	112	215	0.914636804	724	
G10405 034	-	1.08	833	5.8	0.03	4.2	0.09	1.09	1.8	0.45	11	3.8	56	25	135	382	89	12054	464	670	0.03361381	577	
G10405 035	-	2.48	2237	2.9	0.07	25	0.28	4.61	8.7	1.8	51	17	218	81	361	650	124	10592	157	211	0.3939897	632	
G10405 036	-	3.31	2141	8.8	0.36	2.2	0.17	1.49	2.0	0.48	16	8.6	151	69	381	954	200	19038	88	5406	0.520450964	652	
G10405 038	-	7.11	2866	12.6	0.32	1.7	0.19	1.67	2.4	0.53	23	13	214	90	487	1407	294	18771	87	8019	0.851755007	712	
G10405 039	-	bdl	2775	10.8	0.54	3.6	0.23	1.96	2.7	0.65	22	12	199	90	475	1114	226	17361	127	6488	bdl		
G10405 042	-	6.39	2648	10.9	0.95	3.2	0.42	3.14	3.5	0.90	27	13	200	82	427	1257	294	17651	62	5585	0.805750805	703	
G10405 044	-	bdl	2250	7.2	0.03	1.7	0.05	0.60	1.4	0.12	17	9.5	165	76	373	692	132	16700	84	4382	bdl		
G10405 045	-	14.8	3148	27	225	470	56	239	79	2	126	31	318	102	463	1321	306	30804	133	3509	1.17033651	777	
G10405 046	-	1.45	464	1.87	0.02	0.21	0.04	0.48	0.34	0.05	3	1.66	32	15	87	282	65	23831	9	1632	0.162698655	596	
G10405 048	-	4.35	2526	10.13	0.05	2.01	0.05	0.65	1.72	0.12	17	10	178	82	442	1090	225	18507	117	6442	0.638377542	673	
G10405 050	-	4.67	1458	4.56	0.02	1.67	0.05	0.56	0.84	0.08	9	5.7	104	48	247	528	104	17076	41	2762	0.669399534	678	
G10405 051	-	3.62	1907	6.3	0.59	2.3	0.34	2.7	2.2	0.42	15	8.1	136	61	325	801	170	19009	60	3765	0.558921999	659	
G10405 052	-	4.34	6890	20	18	71	5.4	30	24	12	125	47	636	227	972	1906	366	18246	243	8981	0.637967392	673	
G10405 053	-	bdl	1255	6.5	0.21	1.22	0.10	1.10	1.12	0.24	9.1	4.8	85	40	227	642	139	19232	33	2862	bdl		

Trace and REE elemental concentrations in ppm
 Elemental concentrations are calibrated relative to the analysis of NIST 610 trace element standard glass
 Zircon crystallization temperature calculated according to the method and formulas of Watson et al. (2006)
 bdl: below detection limits
 -: not determined

TABLE II | U-Pb data, Chuacús Complex, Central Guatemala

Sample number	U ¹ (ppm)	Th ¹ (ppm)	Th/U	CORRECTED RATIOS ²				CORRECTED RATIOS ³				CORRECTED AGES (Ma)				% disc					
				²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁶ Pb/ ²³⁸ U ±1σ	²⁰⁸ Pb/ ²³² Th ±1σ	Rho	²⁰⁶ Pb/ ²³⁸ U ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	Best age (Ma) ±1σ									
GI0309 Metapelite: 15°01'18.3" 90°42'53.9"																					
GI0309 001	513	169	0.33	0.07275	0.0007	1.6085	0.0169	0.16054	0.0007	0.0485	0.0006	0.4	974	7	1007	19	960	4	1.4		
GI0309 002	608	114	0.19	0.07076	0.0007	1.39806	0.0176	0.14329	0.0009	0.04344	0.0003	0.6	863	5	884	17	863	5	2.8		
GI0309 003	711	84	0.12	0.07209	0.0007	1.3886	0.0156	0.13987	0.0009	0.04342	0.0006	0.6	844	4	988	17	844	5	4.5		
GI0309 004	449	63	0.42	0.07434	0.0013	1.183	0.0215	0.11649	0.0007	0.05555	0.0008	0.3	710	4	1061	34	740	4	40.9		
GI0309 005	206	68	0.33	0.07429	0.001	1.73	0.0279	0.16889	0.001	0.05092	0.0003	0.8	1006	6	1049	27	1049	27	4.1		
GI0309 006	1050	132	0.13	0.06966	0.0007	1.18858	0.0211	0.12374	0.0015	0.03759	0.0005	0.5	752	9	752	21	752	9	5.4		
GI0309 007	612	220	0.36	0.07084	0.0009	1.39745	0.0228	0.14286	0.0009	0.0433	0.0003	0.7	861	5	888	10	861	5	3.0		
GI0309 008	836	131	0.16	0.07082	0.0008	1.34313	0.0255	0.13793	0.0017	0.04183	0.0005	0.7	863	10	956	26	861	10	3.7		
GI0309 009	2026	449	0.07	0.06644	0.0006	0.66559	0.0093	0.07937	0.0004	0.02284	0.0002	0.6	459	2	746	49	459	2	46.4		
GI0309 010	116	46	0.40	0.07292	0.0012	1.54793	0.0288	0.15397	0.0008	0.04652	0.0003	0.3	923	5	954	32	923	5	2.8		
GI0309 011	374	87	0.23	0.07208	0.0008	1.55972	0.0191	0.15693	0.0008	0.04747	0.0002	0.5	940	4	988	20	940	4	1.5		
GI0309 012	528	152	0.29	0.07177	0.0008	1.5127	0.0187	0.15287	0.0007	0.04627	0.0002	0.4	917	4	936	21	917	4	2.0		
GI0309 013	1052	131	0.12	0.07115	0.0007	1.20944	0.0146	0.12329	0.0008	0.03738	0.0002	0.5	749	4	962	19	749	4	7.0		
GI0309 014	1935	240	0.12	0.06482	0.0007	0.7824	0.0109	0.08754	0.0007	0.02683	0.0002	0.7	541	4	769	20	541	4	7.8		
GI0309 015	2333	284	0.13	0.06877	0.0006	0.91113	0.0092	0.09609	0.0004	0.02926	0.0001	0.5	594	2	658	47	594	2	40.2		
GI0309 016	2338	267	0.11	0.06523	0.0006	0.90328	0.0096	0.10044	0.0004	0.03074	0.0002	0.4	617	2	782	19	617	2	5.5		
GI0309 017	500	95	0.19	0.07003	0.0007	1.45129	0.0169	0.1503	0.0007	0.04561	0.0002	0.4	903	4	929	20	903	4	0.8		
GI0309 018	1989	232	0.12	0.06731	0.0006	0.86275	0.0093	0.09296	0.0004	0.02837	0.0001	0.5	573	3	847	18	573	3	9.3		
GI0309 019	4373	448	0.09	0.06569	0.0007	0.76994	0.0085	0.07762	0.0005	0.02377	0.0002	0.6	482	3	747	48	482	3	40.9		
GI0309 020	1911	441	0.23	0.06266	0.0008	0.87439	0.0125	0.1012	0.0005	0.03111	0.0002	0.4	621	3	638	7	697	25	621	3	2.7
GI0309 021	2845	526	0.18	0.06539	0.0007	0.77848	0.0115	0.08635	0.0008	0.02644	0.0002	0.7	534	4	585	7	787	20	534	4	8.7
GI0309 022	344	110	0.32	0.06807	0.0009	1.24584	0.0197	0.13273	0.0008	0.04041	0.0003	0.5	803	5	822	9	803	5	2.3		
GI0309 023	211	86	0.41	0.07151	0.0012	1.50035	0.0332	0.15217	0.0015	0.04607	0.0005	0.6	913	9	931	13	913	9	1.9		
GI0309 024	2263	230	0.10	0.06597	0.0007	0.80816	0.0095	0.08885	0.0005	0.02718	0.0002	0.5	549	3	805	20	549	3	8.7		
GI0309 025	324	53	0.16	0.07077	0.0008	1.4934	0.0189	0.15304	0.0008	0.04639	0.0002	0.4	918	4	928	8	918	4	1.1		
GI0309 026	223	85	0.38	0.07276	0.0009	1.237	0.0163	0.12352	0.0007	0.05038	0.0002	0.4	751	4	818	7	751	4	8.2		
GI0309 027	1059	244	0.23	0.06491	0.0009	1.02931	0.0283	0.11502	0.0021	0.03521	0.0006	0.9	702	12	719	14	702	12	2.4		
GI0309 028	377	174	0.46	0.07313	0.0009	1.68506	0.0289	0.16713	0.0014	0.05047	0.0004	0.6	996	7	1003	11	1017	24	2.1		
GI0309 029	2175	174	0.08	0.0709	0.0006	1.20841	0.0117	0.12362	0.0005	0.03749	0.0002	0.4	751	3	954	17	751	3	6.6		
GI0309 030	2165	611	0.28	0.06682	0.0007	1.17028	0.0144	0.12703	0.0006	0.03876	0.0002	0.6	771	4	787	7	771	4	2.0		
GI0309 031	1163	194	0.17	0.06882	0.0007	1.18335	0.0143	0.12507	0.0008	0.03806	0.0002	0.6	760	4	793	7	760	4	4.2		
GI0309 032	4449	81	0.06	0.06498	0.0007	0.63183	0.0106	0.07045	0.0008	0.02163	0.0003	0.8	439	5	447	24	439	5	11.7		
GI0309 033	1192	262	0.22	0.05999	0.0012	0.77025	0.0183	0.09312	0.0008	0.02876	0.0003	0.6	574	5	580	10	603	40	574	5	1.0
GI0309 034	1298	73	0.06	0.05975	0.0008	0.41601	0.0071	0.05005	0.0005	0.01565	0.0002	0.6	318	3	353	5	318	3	9.9		
GI0309 035	168	102	0.61	0.07596	0.0008	1.8654	0.0237	0.1784	0.0011	0.05643	0.0008	0.5	1058	6	1069	8	1094	21	3.3		
GI0309 036	770	173	0.22	0.07233	0.0007	1.6352	0.0167	0.16423	0.0007	0.04907	0.0006	0.4	980	4	984	6	985	18	980	4	0.4
GI0309 037	665	127	0.19	0.07154	0.0007	1.5825	0.016	0.16056	0.0006	0.04891	0.0006	0.4	960	4	963	6	973	18	960	4	0.3
GI0309 038	1000	160	0.16	0.06974	0.0007	1.21279	0.0129	0.12612	0.0005	0.03831	0.0002	0.4	766	3	806	6	921	18	766	3	5.0
GI0309 039	1099	256	0.23	0.07194	0.0007	1.36868	0.0152	0.13799	0.0007	0.04177	0.0002	0.5	833	4	876	7	984	18	833	4	4.9
GI0309 040	1284	266	0.21	0.06968	0.0007	1.9079	0.0147	0.12395	0.0008	0.03785	0.0002	0.6	753	4	796	7	919	19	753	4	5.4
GI0309 041	3076	447	0.15	0.06126	0.0007	0.55771	0.0109	0.06603	0.0009	0.02037	0.0003	0.8	412	5	450	7	648	22	412	5	8.4
GI0309 042	2338	249	0.11	0.06917	0.0006	1.15717	0.0124	0.12133	0.0002	0.03689	0.0002	0.5	738	4	781	6	738	4	5.5		
GI0309 043	2621	442	0.17	0.06386	0.0007	0.63221	0.0103	0.0957	0.0005	0.02936	0.0002	0.5	589	3	623	21	589	3	5.2		
GI0309 044	1821	166	0.09	0.06447	0.0006	0.87321	0.0076	0.07574	0.0003	0.02324	0.0001	0.4	471	2	575	21	471	2	9.9		
GI0309 045	239	132	0.56	0.06915	0.0011	1.37892	0.025	0.14462	0.0008	0.04385	0.0002	0.4	871	5	880	11	903	30	871	5	1.0
GI0309 046	1195	344	0.29	0.0632	0.0008	0.85747	0.0134	0.09839	0.0006	0.03022	0.0002	0.6	605	4	629	7	715	25	605	4	3.8
GI0309 047	1619	164	0.10	0.07074	0.0006	1.183	0.012	0.12146	0.0006	0.03703	0.0005	0.5	739	3	793	6	950	3	6.8		
GI0309 048	901	137	0.15	0.07107	0.0007	1.33183	0.0162	0.13359	0.0009	0.04119	0.0003	0.5	821	5	860	7	959	20	821	5	4.5
GI0309 049	1721	340	0.20	0.07043	0.0007	1.29521	0.0147	0.13338	0.0007	0.04046	0.0002	0.5	807	4	844	6	941	19	807	4	4.4
GI0309 050	128	44	0.34	0.07231	0.0011	1.58898	0.0289	0.15937	0.001	0.04819	0.0003	0.4	953	6	995	32	953	6	1.3		
GI0309 051	1976	165	0.08	0.05361	0.0007	0.37253	0.0057	0.0504	0.0003	0.01578	0.0002	0.5	317	2	322	4	355	29	317	2	1.6
GI0309 052	219	43	0.20	0.05838	0.0013	0.70152	0.0184	0.08715	0.0008	0.027	0.0004	0.4	540	11	544	49	539	4	0.2		
GI0309 053	496	98	0.50	0.07673	0.0008	1.47663	0.0231	0.14673	0.0012	0.05848	0.0008	0.6	997	7	1033	8	1144	22</			

TABLE II | Continued

Sample number	U ¹ (ppm)	Th ¹ (ppm)	Th/U	CORRECTED RATIOS ²				CORRECTED AGES (Ma)				% disc.											
				²⁰⁷ Pb/ ²⁰⁶ Pb	$\pm 1\sigma$	²⁰⁷ Pb/ ²³⁵ U	$\pm 1\sigma$	²⁰⁶ Pb/ ²³⁸ U	$\pm 1\sigma$	²⁰⁷ Pb/ ²³⁵ U	$\pm 1\sigma$		²⁰⁶ Pb/ ²³⁸ U	$\pm 1\sigma$	Beest age (Ma)	$\pm 1\sigma$							
G10320 018	49	24	0.49	0.05201	±0.0017	0.25288	0.0086	0.03528	0.0003	0.2	224	2	229	7	286	70	224	2	2.2				
G10320 019	74	50	0.67	0.06444	±0.0049	0.27976	0.0086	0.03414	0.0003	0.3	200	2	250	7	746	60	200	2	20.0				
G10320 020	147	101	0.69	0.05243	±0.0011	0.26328	0.0055	0.03647	0.0002	0.3	231	1	237	4	304	42	231	1	2.5				
G10320 021	105	115	1.10	0.05131	±0.0013	0.25308	0.0088	0.03583	0.0003	0.3	227	2	229	6	255	55	227	2	0.9				
G10320 022	89	33	0.38	0.05156	±0.0013	0.25543	0.0066	0.03599	0.0003	0.3	228	2	231	5	266	53	228	2	1.3				
G10320 023	86	62	0.65	0.05087	±0.0016	0.24895	0.0087	0.03549	0.0003	0.3	225	2	226	7	235	67	225	2	0.4				
G10320 024	65	65	1.00	0.06372	±0.0028	0.283	0.014	0.03221	0.0004	0.3	204	2	253	11	732	91	204	2	19.4				
G10320 025	74	68	0.92	0.05165	±0.0024	0.26023	0.0124	0.0366	0.0004	0.3	232	3	235	10	270	104	232	3	1.3				
G10320 026	101	74	0.74	0.05144	±0.0011	0.25444	0.006	0.03591	0.0003	0.3	227	2	230	5	261	47	227	2	1.3				
G10320 027	67	79	1.17	0.0527	±0.0016	0.26	0.0083	0.03585	0.0003	0.3	227	2	235	7	316	65	227	2	3.4				
G10320 028	74	70	0.95	0.05382	±0.0017	0.26713	0.0085	0.03596	0.0003	0.3	228	2	240	7	364	69	228	2	5.0				
G10320 029	37	31	0.84	0.05355	±0.0026	0.24643	0.0133	0.03338	0.0004	0.3	212	2	224	11	352	107	212	2	5.4				
G10320 030	52	48	0.91	0.05399	±0.0026	0.25513	0.0135	0.03427	0.0004	0.3	217	2	231	11	371	106	217	2	5.4				
G10320 031	441	482	1.30	0.05767	±0.0046	0.25175	0.0073	0.03459	0.0003	0.3	200	4	228	6	512	60	200	4	12.3				
G10320 032	74	86	1.16	0.0553	±0.0018	0.26207	0.0087	0.03457	0.0003	0.3	219	2	236	7	424	70	219	2	7.2				
G10320 033	63	55	0.86	0.05121	±0.0015	0.25297	0.0078	0.03599	0.0003	0.3	228	2	229	6	250	68	228	2	0.4				
G10320 034	85	40	0.47	0.05283	±0.0016	0.25602	0.0079	0.03528	0.0003	0.3	224	2	231	6	322	67	224	2	3.0				
G10320 035	60	57	0.95	0.05174	±0.0016	0.23192	0.0072	0.03259	0.0003	0.3	207	2	212	6	274	68	207	2	2.4				
G10320 036	195	175	0.90	0.05192	±0.0009	0.24781	0.0047	0.03461	0.0002	0.4	219	1	225	4	282	40	219	1	2.7				
G10320 037	88	60	0.68	0.05301	±0.0015	0.26491	0.0079	0.03625	0.0004	0.4	230	2	239	6	329	62	230	2	3.8				
G10320 038	180	299	1.66	0.05174	±0.0009	0.24858	0.0047	0.03491	0.0002	0.3	221	1	225	4	274	41	221	1	1.8				
G10320 039	240	260	1.09	0.05072	±0.0008	0.24573	0.0042	0.03517	0.0002	0.3	223	1	223	3	228	36	223	1	0.0				
G10320 040	247	303	1.23	0.052	±0.0008	0.25266	0.0043	0.03528	0.0002	0.3	224	1	229	3	285	36	224	1	2.2				
G10320 041	185	120	0.65	0.05303	±0.001	0.26095	0.0052	0.03571	0.0002	0.3	226	1	235	4	330	43	226	1	3.8				
G10320 042	104	140	1.34	0.05217	±0.0015	0.25558	0.0075	0.03561	0.0003	0.3	226	2	231	6	293	63	226	2	2.2				
G10320 043	203	176	0.87	0.0521	±0.0008	0.25456	0.0044	0.03549	0.0002	0.4	225	1	230	4	290	36	225	1	2.2				
G10320 044	66	66	1.00	0.05325	±0.0016	0.25415	0.0079	0.03475	0.0003	0.2	220	2	230	6	339	67	220	2	4.3				
G10321, Agua Caliente River felsic orthogneiss, 14°56'07.8" 90°30'10.9"																							
G10321 006	431	218	0.51	0.05126	±0.0017	0.24029	0.0088	0.03395	0.0004	0.4	215	3	219	7	253	72	215	3	1.8				
G10321 007	179	95	0.53	0.05188	±0.0032	0.20801	0.0136	0.02908	0.0004	0.3	185	3	192	11	280	127	185	3	3.6				
G10321 008	488	460	0.85	0.06464	±0.0036	0.24863	0.0133	0.02578	0.0005	0.3	464	3	201	11	662	115	464	3	18.4				
G10321 009	74	46	0.62	0.05423	±0.0004	0.03277	0.0006	0.03277	0.0006	0.2	208	3	223	16	381	153	208	3	6.7				
G10321 010	937	726	0.78	0.05064	±0.0013	0.23386	0.0064	0.03344	0.0004	0.4	212	2	213	5	224	53	212	2	0.5				
G10321 011	479	462	0.90	0.05747	±0.0034	0.27443	0.0175	0.03463	0.0004	0.2	219	2	246	14	510	119	219	2	11.0				
G10321 013	193	155	0.81	0.05688	±0.0022	0.27085	0.0106	0.03462	0.0004	0.3	219	2	243	9	487	77	219	2	9.9				
G10321 014	264	494	0.73	0.05866	±0.0025	0.29498	0.0141	0.03647	0.0003	0.3	234	2	262	11	554	87	234	2	11.8				
G10321 015	416	86	0.75	0.06322	±0.0037	0.32066	0.0189	0.03664	0.0004	0.2	232	2	288	15	716	114	232	2	18.0				
G10321 016	170	176	1.04	0.05536	±0.0032	0.28553	0.0166	0.03745	0.0004	0.2	237	3	255	13	427	118	237	3	7.1				
G10321 017	472	52	0.29	0.06761	±0.0061	0.317	0.025	0.034	0.0004	0.3	246	2	290	19	857	144	246	2	22.9				
G10321 019	162	128	0.79	0.05307	±0.0022	0.26791	0.0119	0.03634	0.0005	0.3	230	3	241	9	332	88	230	3	4.6				
G10321 020	263	185	0.70	0.0561	±0.0031	0.30224	0.0178	0.03808	0.0003	0.2	247	2	268	14	456	113	247	2	7.8				
G10321 021	635	574	0.90	0.05042	±0.0014	0.23915	0.0071	0.03437	0.0003	0.3	218	2	218	6	214	60	218	2	0.0				
G10321 022	747	666	0.89	0.05077	±0.001	0.23847	0.0053	0.03399	0.0003	0.3	215	2	217	4	230	43	215	2	0.9				
G10321 024	217	172	0.82	0.05987	±0.0047	0.29938	0.0087	0.03582	0.0003	0.3	227	2	262	7	588	56	227	2	13.4				
G10321 025	217	195	0.90	0.05706	±0.0019	0.29318	0.0104	0.03711	0.0004	0.3	235	2	264	8	494	69	235	2	10.0				
G10321 026	443	92	0.64	0.05845	±0.0027	0.27575	0.013	0.03424	0.0004	0.2	247	2	247	40	547	69	247	2	12.1				
G10321 027	121	91	0.76	0.05678	±0.0023	0.30485	0.0129	0.03885	0.0004	0.3	246	3	270	10	483	83	246	3	8.9				
G10321 028	213	163	0.76	0.06631	±0.0057	0.33287	0.0306	0.03641	0.0005	0.2	234	3	292	23	816	167	234	3	20.9				
G10321 030	167	66	0.39	0.04961	±0.0053	0.22677	0.0254	0.03314	0.0005	0.2	210	3	207	21	177	215	210	3	1.4				
G10321 031	86	48	0.56	0.0671	±0.0042	0.32966	0.021	0.03567	0.0005	0.2	226	3	288	16	841	124	226	3	21.8				
G10321 032	147	160	0.68	0.06806	±0.003	0.28733	0.0155	0.03586	0.0005	0.2	227	3	266	12	532	107	227	3	11.3				
G10321 033	1942	147	0.08	0.07937	±0.0013	1.96167	0.0343	0.17926	0.0012	0.4	1063	7	1102	12	1181	30	1181	30	10.0				
G10321 034	464	297	0.64	0.06082	±0.0013	0.48084	0.0114	0.05725	0.0005	0.4	359	3	399	8	633	45	359	3	10.0				
G10321 036	379	154	0.40	0.07518	±0.0044	1.6405	0.0289	0.15893	0.0009	0.3	947	6	986	13	1073	36	1073	36	11.7				
G10321 037	1453	291	0.20	0.07227	±0.0012	1.5936	0.0283	0.15958	0.0012	0.4	954	7	988	11	994	31	984	7	1.4				
G10321 038	201	197	0.98	0.06544	±0.0048	0.34677	0.0094	0.03593	0.0004	0.3	224	2	279	7	788	55	224	2	19.7				
G10321 039	486	248	1.17	0.06744	±0.0061	0.08744	0.0324	0.03324	0.0007	0.4	211	4	342	23	1370	130	211	4	38.3				
G10321 040	213	41	0.05																				

TABLE II | Continued

Sample number	U ¹ (ppm)	Th ¹ (ppm)	Th/U	CORRECTED RATIOS ²				CORRECTED RATIOS ¹										Best age (Ma)	±1σ	±1σ	% disc		
				²⁰⁷ Pb/ ²⁰⁹ Pb	²⁰⁷ Pb/ ²³⁵ U	²⁰⁷ Pb/ ²³⁸ U	²⁰⁸ Pb/ ²³² Th	Rho	²⁰⁶ Pb/ ²³⁸ U	²⁰⁷ Pb/ ²³⁸ U	²⁰⁷ Pb/ ²³⁵ U	²⁰⁸ Pb/ ²³² Th	²⁰⁶ Pb/ ²³⁸ U	²⁰⁷ Pb/ ²³⁸ U	²⁰⁷ Pb/ ²³⁵ U	²⁰⁸ Pb/ ²³² Th	²⁰⁶ Pb/ ²³⁸ U					²⁰⁷ Pb/ ²³⁸ U	²⁰⁷ Pb/ ²³⁵ U
310332-050	46	32	0.69	0.13981	1.2565	0.0992	0.0606	0.06512	0.0009	0.10951	0.0009	0.03362	0.0003	0.3	670	5	680	23	713	92	670	5	1.5
310332-051	47	28	0.59	0.06315	0.95347	0.0448	0.11917	0.0001	0.04448	0.0006	0.4	840	5	865	11	929	35	840	5	1071	65	2.9	
310332-052	60	70	1.18	0.07002	1.3441	0.0247	0.17886	0.0001	0.05342	0.0011	0.4	1050	12	1057	22	1071	65	1071	65	1071	65	2.0	
310332-053	7	7	0.98	0.07508	1.8334	0.0627	0.14967	0.0001	0.04448	0.0006	0.4	965	7	1044	77	1308	244	1308	244	1308	244	3.6	
310332-054	26	48	0.72	0.04468	1.7452	0.2063	0.17968	0.0009	0.05463	0.0008	0.5	1065	5	1079	7	1105	20	1105	20	1105	20	3.6	
310332-055	211	61	0.29	0.07639	1.895	0.0214	0.14967	0.0001	0.04994	0.0003	0.5	988	6	1009	11	1053	29	1053	29	1053	29	6.2	
310332-056	48	13	0.27	0.07443	1.70055	0.0292	0.1657	0.0011	0.05174	0.0006	0.4	1043	4	1054	7	1071	19	1071	19	1071	19	2.6	
310332-057	492	237	0.48	0.07508	1.8209	0.0188	0.17568	0.0008	0.05259	0.0007	0.5	1018	6	1034	7	1067	19	1067	19	1067	19	4.6	
310332-058	347	134	0.39	0.07494	1.7691	0.02	0.17103	0.0001	0.05417	0.0008	0.4	1053	5	1074	8	1123	26	1123	26	1123	26	-0.2	
310332-059	189	47	0.25	0.07436	1.8223	0.0219	0.17749	0.0009	0.05481	0.0009	0.4	1050	5	1074	9	1123	26	1123	26	1123	26	6.5	
310332-060	77	42	0.54	0.07708	1.8798	0.0265	0.17884	0.0001	0.05481	0.0009	0.4	1050	5	1074	9	1123	26	1123	26	1123	26	6.5	
310332-061	578	53	0.09	0.05113	0.26071	0.0124	0.03711	0.0005	0.02728	0.003	0.3	235	3	235	10	247	95	235	3	235	3	0.0	
310332-062	479	42	0.09	0.05066	0.24796	0.0114	0.03546	0.0005	0.0232	0.0028	0.2	225	3	225	9	225	91	225	3	225	3	0.0	
310332-063	605	57	0.09	0.04994	0.25055	0.0131	0.03651	0.0004	0.01321	0.0028	0.3	231	3	227	11	192	106	231	3	192	106	3	
310332-064	1494	110	0.07	0.05484	0.02640	0.0098	0.03496	0.0003	0.02604	0.0034	0.3	222	2	238	8	406	73	222	2	406	73	6.7	
310332-065	1888	374	0.20	0.05049	0.24094	0.0025	0.03461	0.0003	0.08597	0.0059	0.3	219	2	219	2	218	3	219	2	218	3	0.0	
310332-066	87	34	0.39	0.05049	0.24103	0.0027	0.03462	0.0004	0.04323	0.0049	0.3	219	2	219	2	218	3	219	2	218	3	0.0	
310332-067	275	63	0.23	0.05627	0.27975	0.0124	0.03659	0.0005	0.02926	0.0044	0.3	232	3	250	10	463	84	232	3	463	84	7.2	
310341, orthogneiss, 15°00'22.5" 90°14'24.4"																							
310341-001	171	146	0.85	0.05696	0.57685	0.0093	0.07356	0.0004	0.02256	0.0003	0.4	458	3	462	6	490	32	458	3	490	32	0.9	
310341-002	121	90	0.75	0.05805	0.62159	0.0106	0.07786	0.0004	0.02405	0.0003	0.3	483	3	491	7	532	35	483	3	532	35	1.6	
310341-003	86	43	0.50	0.0726	1.6717	0.0235	0.16698	0.0009	0.04928	0.0007	0.4	995	5	998	9	1003	26	998	9	1003	26	0.3	
310341-004	509	398	0.78	0.05594	0.52663	0.0059	0.06884	0.0004	0.02015	0.0002	0.5	427	2	430	4	450	22	427	2	450	22	0.7	
310341-005	825	1760	2.13	0.06273	1.60464	0.0141	0.0698	0.0014	0.02303	0.0012	0.9	435	8	480	9	689	25	435	8	689	25	9.4	
310341-006	608	137	0.22	0.05622	0.56093	0.0061	0.07241	0.0003	0.0233	0.0003	0.4	451	2	452	4	461	21	451	2	461	21	0.2	
310341-007	165	130	0.79	0.05993	0.63942	0.0192	0.07739	0.0007	0.0239	0.0002	0.4	481	4	502	12	601	53	481	4	601	53	4.2	
310341-008	660	606	0.92	0.05813	0.54479	0.0075	0.06804	0.0006	0.02017	0.0004	0.6	424	3	442	5	535	23	424	3	535	23	4.1	
310341-009	326	279	0.86	0.05944	0.61262	0.0088	0.07485	0.0004	0.02414	0.0003	0.4	465	3	485	6	583	27	465	3	583	27	4.1	
310341-010	170	120	1.23	0.05707	0.59089	0.0114	0.07524	0.0005	0.02335	0.0003	0.3	469	3	471	5	494	39	469	3	494	39	0.6	
310341-011	333	224	0.67	0.05674	0.58979	0.0073	0.07552	0.0004	0.02228	0.0003	0.5	468	3	471	5	481	24	468	3	481	24	0.4	
310341-012	429	241	0.56	0.05584	0.57358	0.0064	0.07459	0.0004	0.02316	0.0003	0.5	464	2	460	4	446	22	464	2	446	22	-0.9	
310341-013	84	103	1.22	0.05721	0.6047	0.0104	0.07879	0.0005	0.02416	0.0003	0.4	477	3	480	7	500	35	477	3	500	35	0.6	
310341-014	96	94	0.94	0.07682	1.7665	0.0175	0.07284	0.0006	0.02665	0.0004	0.3	463	3	574	49	1168	44	463	3	1168	44	24.1	
310341-015	514	339	0.66	0.05649	0.56210	0.0062	0.07225	0.0004	0.02185	0.0002	0.5	450	2	453	4	472	22	450	2	472	22	0.7	
310341-016	653	941	1.44	0.05719	0.56314	0.0065	0.07162	0.0004	0.02191	0.0003	0.5	446	3	454	4	499	21	446	3	499	21	4.6	
310341-017	160	209	1.31	0.05624	0.59385	0.0092	0.07674	0.0005	0.02436	0.0003	0.4	477	3	473	6	462	31	477	3	462	31	-0.8	
310341-018	396	740	1.87	0.05658	0.55058	0.0071	0.07072	0.0003	0.02142	0.0002	0.4	440	2	445	5	475	26	440	2	475	26	1.1	
310341-019	463	585	1.26	0.05709	0.56762	0.0064	0.07227	0.0004	0.02268	0.0003	0.5	450	2	456	4	495	22	450	2	495	22	1.3	
310341-020	1187	1065	0.90	0.05559	0.40887	0.0079	0.05314	0.0009	0.01643	0.0002	0.9	334	6	348	6	448	20	334	6	448	20	6.0	
310341-021	215	89	0.41	0.05773	0.58327	0.0081	0.07341	0.0004	0.02287	0.0003	0.4	457	2	467	5	520	28	457	2	520	28	2.1	
310341-022	780	461	0.59	0.057	0.56357	0.0062	0.07181	0.0004	0.02246	0.0003	0.5	447	2	454	4	492	21	447	2	492	21	1.5	
310341-023	227	163	0.72	0.05571	0.5853	0.0087	0.07639	0.0004	0.02386	0.0003	0.4	475	2	468	6	441	31	475	2	441	31	-1.5	
310341-024	429	137	0.32	0.09328	3.2453	0.0315	0.25199	0.0011	0.02075	0.0008	0.5	1449	6	1468	8	1494	16	1468	8	1494	16	3.0	
310341-025	429	160	0.37	0.05931	0.001	0.74541	0.015	0.09115	0.0005	0.02554	0.0003	0.4	291	3	313	7	530	20	291	3	7.0		
310341-026	226	267	1.18	0.05684	0.59058	0.0089	0.07545	0.0004	0.02336	0.0003	0.4	469	3	471	6	485	31	469	3	485	31	0.4	
310341-027	120	150	1.26	0.05832	0.60226	0.0116	0.07504	0.0005	0.02407	0.0003	0.4	466	3	479	7	542	39	466	3	542	39	2.7	
310341-028	388	476	1.23	0.0578	0.57568	0.0076	0.07238	0.0004	0.02303	0.0003	0.4	450	2	462	5	522	24	450	2	522	24	2.6	
310341-029	364	216	0.59	0.05687	0.57738	0.0076	0.0738	0.0004	0.02344	0.0003	0.4	459	2	463	5	487	26	459	2	487	26	0.9	
310341-030	214	94	0.44	0.05744	0.56823	0.0088	0.07187	0.0005	0.02234	0.0003	0.4	447	3	457	6	509	30	447	3	509	30	2.2	
310341-031	305	258	0.85	0.06192	0.60361	0.014	0.07444	0.0013	0.02034	0.0012	0.8	463	8	498	9	671	30	463	8	671	30	7.0	
310341-032	331	130	0.39	0.05708	0.59749	0.0073	0.07595	0.0004	0.02361	0.0003	0.4	472	2	476	5	495	24	472	2	495	24	0.8	
310341-033	117	90	0.77	0.058	0.36057	0.0092	0.04619	0.0005	0.02554	0.0003	0.4	291	3	313	7	530	20	291	3	530	20	7.0	
310341-034	144	86	0.59	0.0571	0.57709	0.0082	0.07341	0.0004	0.02258	0.0003	0.4	457	2	463	5	495	28	457	2	495	28	1.3	
310341-035	94	257	2.74	0.05771	0.001	0.57238	0.0104	0.07208	0.0005	0.02227	0.0003	0.4	449	3	460	7	519	37	449	3	519	37	2.4
310341-036	105	224	2.13	0.05773	0.0009	0.59391	0.0096	0.0747	0.0004	0.02229	0.0003	0.4	464	3	473	6	520	33	464	3	520	33</	

TABLE II | Continued

Sample number	U ¹ (ppm)	Th ¹ (ppm)	Th/U	CORRECTED RATIOS ²										CORRECTED AGES (Ma)										% disc
				²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁶ Pb/ ²³⁵ U ±1σ	²⁰⁸ Pb/ ²³² Th ±1σ	²⁰⁸ Pb/ ²³⁵ U ±1σ	Rho	²⁰⁶ Pb/ ²³⁸ U ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁶ Pb/ ²³⁵ U ±1σ	Best age (Ma) ±1σ	±1σ	±1σ	±1σ	±1σ	±1σ	±1σ	±1σ	±1σ	±1σ		
GI0341 046	620	355	0.57	0.05663	0.0006	0.57147	0.0063	0.07329	0.0004	0.02283	0.0003	0.4	456	2	459	4	477	22	456	20	0.4			
GI0341 047	223	378	1.70	0.06149	0.0008	0.58726	0.0089	0.06937	0.0005	0.02113	0.0004	0.5	432	3	469	6	656	27	432	3	7.9			
GI0341 048	550	345	0.63	0.05665	0.0006	0.57569	0.0069	0.07386	0.0005	0.02223	0.0003	0.4	459	2	462	4	478	24	459	2	0.6			
GI0341 049	222	293	1.32	0.05911	0.0009	0.58091	0.0096	0.07134	0.0003	0.02469	0.0003	0.4	444	3	465	6	571	32	444	3	4.5			
GI0341 050	195	47	0.24	0.05777	0.0009	0.56646	0.0092	0.07126	0.0004	0.02264	0.0004	0.4	444	3	456	6	521	32	444	3	2.6			
GI03113 pegmatite dike. 15°03'24" 90°23'36.8"																								
GI03113 001	256	80	0.31	0.07737	0.0009	2.0375	0.0243	0.19089	0.0009	0.05905	0.0009	0.4	1126	5	1128	8	1131	20	1131	20	0.4			
GI03113 002	300	92	0.31	0.07233	0.001	1.59942	0.0261	0.16037	0.0008	0.04848	0.0002	0.4	959	4	970	10	995	28	959	4	1.1			
GI03113 003	1051	122	0.12	0.06444	0.0013	1.00002	0.0223	0.11255	0.0006	0.03447	0.0005	0.5	688	4	704	11	756	41	688	4	2.3			
GI03113 004	496	49	0.18	0.06149	0.0008	0.58726	0.0089	0.06937	0.0005	0.02113	0.0004	0.5	432	3	469	6	656	27	432	3	7.9			
GI03113 005	277	102	0.37	0.07603	0.0008	1.8186	0.0202	0.17351	0.0009	0.05358	0.0008	0.3	1031	5	1052	7	1096	19	1031	5	5.9			
GI03113 006	486	79	0.16	0.07096	0.0016	1.41842	0.0354	0.14498	0.0009	0.04393	0.0004	0.3	873	5	897	15	956	45	873	5	2.7			
GI03113 007	245	86	0.35	0.07668	0.0008	1.9842	0.0214	0.18765	0.0008	0.05568	0.0008	0.4	1109	5	1110	7	1113	19	1113	19	0.4			
GI03113 008	866	445	0.17	0.07233	0.001	1.59942	0.0261	0.11255	0.0006	0.03447	0.0005	0.5	688	4	704	11	756	41	688	4	2.3			
GI03113 009	426	54	0.43	0.06395	0.0009	0.63922	0.0222	0.06395	0.0009	0.02222	0.0003	0.4	947	4	1059	8	1291	20	1059	8	43.0			
GI03113 010	386	102	0.27	0.07497	0.0005	1.6892	0.0183	0.16336	0.0008	0.0484	0.0007	0.4	975	4	1004	7	1088	19	1088	19	8.7			
GI03113 011	479	35	0.20	0.07622	0.0005	1.77426	0.0238	0.16746	0.0008	0.05003	0.0007	0.2	998	4	1035	45	1114	428	1114	428	10.4			
GI03113 012	1113	88	0.08	0.05837	0.0011	0.60878	0.0123	0.07564	0.0004	0.02343	0.0003	0.4	470	2	483	8	544	39	470	2	2.7			
GI03113 013	109	32	0.29	0.07575	0.0009	2.0461	0.0264	0.19606	0.0009	0.05932	0.0009	0.4	1154	5	1131	9	1088	23	1088	23	-2.0			
GI03113 014	436	48	0.35	0.06283	0.0019	2.12074	0.0466	0.18657	0.0065	0.05532	0.0042	0.3	1098	36	1166	178	1265	611	1265	611	43.2			
GI03113 015	247	72	0.29	0.07676	0.0009	1.97327	0.0262	0.18645	0.0009	0.056	0.0003	0.4	1102	5	1106	9	1115	22	1115	22	1.2			
GI03113 016	98	32	0.33	0.07246	0.00146	1.64322	0.035	0.16218	0.0033	0.04905	0.0034	0.2	969	18	976	136	991	436	969	18	0.7			
GI03113 017	348	62	0.18	0.07423	0.0012	1.65157	0.0301	0.16137	0.0009	0.04865	0.0004	0.4	964	5	990	12	1048	32	1048	32	8.0			
GI03113 018	383	125	0.45	0.0774	0.0007	2.0198	0.0212	0.18831	0.0009	0.05888	0.0008	0.4	1118	5	1122	7	1132	18	1132	18	11.8			
GI03113 019	117	52	0.33	0.07405	0.0016	1.76258	0.0414	0.17264	0.0009	0.05206	0.0003	0.3	1027	5	1032	15	1043	41	1043	41	1.5			
GI03113 020	115	38	0.33	0.07731	0.0009	2.018	0.0244	0.1894	0.001	0.0553	0.0008	0.4	1118	5	1122	8	1129	21	1129	21	1.0			
GI03113 021	966	149	0.15	0.05726	0.0008	0.61043	0.0101	0.07731	0.0004	0.02401	0.0002	0.4	480	2	484	6	502	31	480	2	0.8			
GI03113 022	240	71	0.30	0.07466	0.0008	1.7564	0.0193	0.17075	0.0008	0.05223	0.0007	0.4	1016	4	1030	7	1059	19	1059	19	4.1			
GI03113 023	113	28	0.25	0.07301	0.002	1.71135	0.0556	0.17	0.0018	0.05134	0.0005	0.5	1012	10	1013	21	1014	55	1014	55	0.4			
GI03113 024	509	67	0.13	0.07067	0.0017	1.34641	0.0351	0.13817	0.0008	0.04188	0.0007	0.4	834	4	866	15	948	47	834	4	3.7			
GI03113 025	368	95	0.26	0.07109	0.0011	1.57555	0.0271	0.15686	0.0008	0.04752	0.0003	0.4	939	5	946	11	960	30	939	5	0.7			
GI03113 026	558	157	0.28	0.07294	0.0008	1.56993	0.0222	0.15522	0.0009	0.04688	0.0003	0.5	930	5	955	9	1012	22	930	5	2.6			
GI03113 027	522	137	0.26	0.07276	0.0013	1.62392	0.0329	0.16187	0.0011	0.04891	0.0004	0.5	967	6	980	13	1007	34	967	6	1.3			
GI03113 028	212	49	0.23	0.07609	0.0008	1.9658	0.0219	0.18741	0.0009	0.05855	0.0008	0.4	1107	5	1104	7	1097	19	1097	19	-0.9			
GI03113 029	208	50	0.24	0.07413	0.0007	1.7159	0.0193	0.16797	0.0009	0.05563	0.0008	0.5	1001	5	1014	7	1045	19	1045	19	4.2			
GI03113 030	120	50	0.42	0.07579	0.0008	1.9	0.0234	0.18189	0.001	0.05603	0.0008	0.5	1077	5	1081	8	1090	21	1090	21	1.2			
GI0405 migmatite Rio Passabien. 15°02'21.4" 89°41'02.4"																								
GI0405 001	91	70	0.77	0.05885	0.0054	0.31666	0.0303	0.03896	0.0004	0.01206	0.0002	0.2	246	3	279	23	265	211	246	3	11.8			
GI0405 002	411	143	0.35	0.05009	0.001	0.28745	0.0062	0.04166	0.0003	0.01242	0.0002	0.4	263	2	257	5	199	43	263	2	-2.3			
GI0405 003	99	34	0.34	0.07037	0.001	1.4623	0.0222	0.15106	0.0009	0.04561	0.0007	0.4	907	5	915	9	939	30	907	5	0.9			
GI0405 004	150	88	0.59	0.05505	0.0011	0.53816	0.0107	0.07111	0.0004	0.02147	0.0003	0.3	443	3	437	7	414	44	443	3	-1.4			
GI0405 005	178	70	0.39	0.05912	0.0026	0.59211	0.0272	0.07263	0.0004	0.02247	0.0002	0.3	452	3	472	17	572	99	452	3	4.2			
GI0405 006	169	69	0.41	0.05746	0.001	0.59155	0.0107	0.07492	0.0005	0.024	0.0004	0.3	466	3	472	7	509	38	466	3	1.3			
GI0405 007	466	181	0.39	0.0553	0.0007	0.56879	0.0074	0.0748	0.0004	0.02359	0.0003	0.4	465	2	457	5	424	27	465	2	-1.8			
GI0405 008	132	50	0.38	0.05595	0.0011	0.5506	0.0115	0.07158	0.0004	0.02195	0.0004	0.3	446	3	445	8	450	46	446	3	-0.2			
GI0405 009	403	262	0.65	0.06292	0.0033	1.8602	0.0402	0.08462	0.0007	0.06321	0.0019	0.4	506	4	1067	14	2486	35	2486	35	52.6			
GI0405 010	264	137	0.52	0.06109	0.0079	0.93669	0.0454	0.0399	0.0005	0.0123	0.0003	0.3	252	3	294	34	257	257	252	3	14.3			
GI0405 011	195	71	0.36	0.05627	0.001	0.55607	0.01	0.07184	0.0004	0.02203	0.0003	0.3	447	3	449	7	463	35	447	3	0.4			
GI0405 012	561	200	0.36	0.05594	0.0006	0.56992	0.0062	0.07232	0.0004	0.02237	0.0003	0.5	450	2	450	4	450	20	450	2	0.0			
GI0405 013	5108	2791	0.55	0.07245	0.0032	0.2156	0.0104	0.02458	0.0002	0.00652	0.0002	0.5	138	1	198	9	999	81	138	1	30.3			
GI0405 014	595	292	0.49	0.05195	0.0007	0.2837	0.0043	0.03964	0.0002	0.01237	0.0002	0.4	251	1	254	3	283	30	251	1	1.2			
GI0405 015	2490	342	0.14	0.05852	0.0018	0.46303	0.0151	0.05739	0.0004	0.01777	0.0002	0.4	360	2	364	11	549	61	360	2	6.7			
GI0405 016	4045	401	0.03	0.06211	0.0009	0.37669	0.006	0.04446	0.0002	0.06733	0.0024	0.3	249	4	324	4	678	30	280	4	13.6			
GI0405 017	570	295	0.52	0.05145	0.0007	0.27898	0.0042	0.03939	0.0002	0.01262	0.0002	0.4	249	1	250	3	261	30	249	1	0.4			
GI0405 018	220	102	0.46	0.05642	0.0012	0.55387	0.0132	0.0712	0.0004	0.02215	0.0001	0.3	443	2	448	9	469	44	443	2	1.1			
GI0405 019	784	519	0.88	0.05312	0.0007	0.27705	0.0039	0.03792	0.0002	0.01195	0.0002	0.4	240	1	248	3	334	27	240	1	3.2			
GI0405 020	705	195	0.28	0.05233	0.0006	0.29051	0.0038	0.04029	0.0002	0.0122	0.0													

TABLE II | Continued

Sample number	U ¹ (ppm)	Th ¹ (ppm)	Th/U	CORRECTED RATIOS ²			CORRECTED RATIOS ³			CORRECTED AGES (Ma)			% disc								
				²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	²⁰⁷ Pb/ ²³⁵ U ±1σ	²⁰⁷ Pb/ ²⁰⁶ Pb ±1σ	Best age (Ma)	±1σ		±1σ							
G10405-026	543	50	0.09	0.05476	0.0007	0.46443	0.0065	0.06143	0.0003	0.0236	0.0004	0.4	384	2	387	5	402	30	384	2	0.8
G10405-027	4241	861	0.20	0.06624	0.0069	0.40377	0.0484	0.04643	0.0007	0.01629	0.0006	0.4	311	4	344	36	676	268	311	4	9.6
G10405-028	386	185	0.48	0.05604	0.0007	0.58126	0.0079	0.07531	0.0005	0.02363	0.0003	0.5	418	3	465	5	454	27	468	3	-0.6
G10405-029	1608	114	0.07	0.05291	0.0005	0.35923	0.0042	0.04921	0.0003	0.02119	0.0003	0.5	310	2	312	3	325	23	310	2	0.6
G10405-030	425	307	0.72	0.05068	0.0008	0.26803	0.0043	0.03844	0.0002	0.01194	0.0002	0.4	243	1	241	3	226	36	243	1	-0.8
G10405-031	188	100	0.53	0.06236	0.001	0.8206	0.0104	0.07239	0.0004	0.02528	0.0004	0.3	451	2	490	6	686	35	451	2	8.0
G10405-032	807	175	0.22	0.06298	0.0082	0.56481	0.0779	0.06504	0.0008	0.01997	0.0013	0.4	406	5	455	61	708	302	406	5	10.8
G10405-033	2871	67	0.02	0.06246	0.0061	0.39816	0.0089	0.03697	0.0003	0.01307	0.0002	0.3	248	2	266	22	660	185	248	2	16.3
G10405-034	590	414	0.70	0.05989	0.0006	0.60858	0.0068	0.07379	0.0004	0.02311	0.0003	0.4	459	2	483	4	600	22	459	2	5.0
G10405-035	186	142	0.77	0.05703	0.001	0.56871	0.0106	0.07249	0.0006	0.02337	0.0004	0.4	451	3	457	7	453	38	451	3	1.3
G10405-036	4761	79	0.02	0.05076	0.0005	0.26191	0.0003	0.03742	0.0002	0.01179	0.0002	0.4	237	1	236	2	230	25	237	1	-0.4
G10405-037	3872	414	0.11	0.06749	0.0284	0.60096	0.1669	0.04682	0.0016	0.01476	0.0009	0.3	313	10	478	106	1371	697	313	10	34.6
G10405-038	5709	113	0.02	0.05097	0.0005	0.27027	0.0029	0.03846	0.0002	0.01211	0.0002	0.5	243	1	243	2	239	22	243	1	0.0
G10405-039	5709	113	0.02	0.05571	0.0041	0.29857	0.0225	0.03887	0.0003	0.01211	0.0028	0.2	246	2	265	18	441	170	246	2	7.2
G10405-040	3298	47	0.01	0.05967	0.0463	0.28217	0.2744	0.03651	0.0028	0.01697	0.3102	0.7	225	18	260	216	562	##	225	18	13.5
G10405-041	6747	112	0.02	0.06819	0.0224	0.39009	0.1226	0.03675	0.0016	0.01698	0.0642	0.5	226	9	284	64	674	695	226	9	23.1
G10405-042	4903	55	0.01	0.05251	0.0011	0.27846	0.0065	0.03946	0.0002	0.01207	0.0011	0.3	243	1	249	5	308	50	243	1	2.4
G10405-043	6407	90	0.01	0.08375	0.0265	0.41899	0.1408	0.03628	0.0014	0.01408	0.0477	0.5	230	9	365	101	1387	681	230	9	35.2
G10405-044	3838	74	0.02	0.05101	0.0005	0.25501	0.0027	0.03628	0.0002	0.01243	0.0002	0.5	230	1	231	2	241	21	230	1	0.4
G10405-045	3095	117	0.04	0.05716	0.0013	0.46073	0.0119	0.05846	0.0004	0.01816	0.001	0.5	366	2	385	8	488	51	366	2	4.9
G10405-046	1431	8	0.01	0.05068	0.0006	0.2418	0.0032	0.03466	0.0002	0.01143	0.0005	0.4	220	1	220	3	226	28	220	1	0.0
G10405-047	7860	101	0.01	0.06913	0.0144	0.36174	0.0782	0.03669	0.0009	0.01421	0.025	0.3	234	6	306	69	663	461	234	6	25.8
G10405-048	5658	105	0.02	0.05148	0.0005	0.27134	0.0037	0.03826	0.0004	0.01608	0.0006	0.8	242	2	244	3	262	21	242	2	0.8
G10405-049	2008	22	0.01	0.07563	0.0013	0.39327	0.0069	0.03809	0.0002	0.01427	0.0169	0.3	244	4	327	5	4082	34	244	4	26.5
G10405-050	2428	36	0.01	0.05111	0.0005	0.26734	0.0029	0.03797	0.0002	0.01593	0.0005	0.5	240	1	241	2	246	22	240	1	0.4
G10405-051	3259	53	0.02	0.05168	0.0008	0.25688	0.0044	0.03606	0.0002	0.01134	0.0009	0.4	228	4	271	36	228	1	271	36	1.7
G10405-052	7897	216	0.03	0.05272	0.0025	0.28703	0.0147	0.03949	0.0003	0.01238	0.0015	0.4	250	2	256	12	317	109	250	2	2.3
G10405-053	2522	29	0.01	0.05129	0.0005	0.26119	0.0003	0.03693	0.0002	0.01162	0.0002	0.4	234	1	236	2	254	23	234	1	0.8
G10405-064	2118	64	0.03	0.06224	0.0066	0.46032	0.0422	0.06626	0.0006	0.01614	0.0023	0.2	330	3	378	30	662	168	330	3	12.7
G10405-065	1822	19	0.01	0.06607	0.0107	0.28922	0.0649	0.03483	0.0006	0.01084	0.0038	0.3	221	4	242	44	455	399	221	4	8.7

¹: U and Th concentrations are calibrated relative to the analysis of NIST 612 trace element standard glass

²: Isotopic ratios are corrected relative to analysis of a standard zircon for mass bias and down-hole fractionation (see text for further explanations and the used zircon standards). The Andersen (2002) common Pb correction method is used.

³: Isotopic ratios are absolute and expressed at ± 1 sigma

Apparent age errors are expressed at ± 1 sigma

Strikethrough analyses were discarded from calculations because >10% <-5% discordant. This criteria was released to >15% for the sample G10315 because of the difficulty to measure 207Pb in such young zircons